

Water resources management and climate change

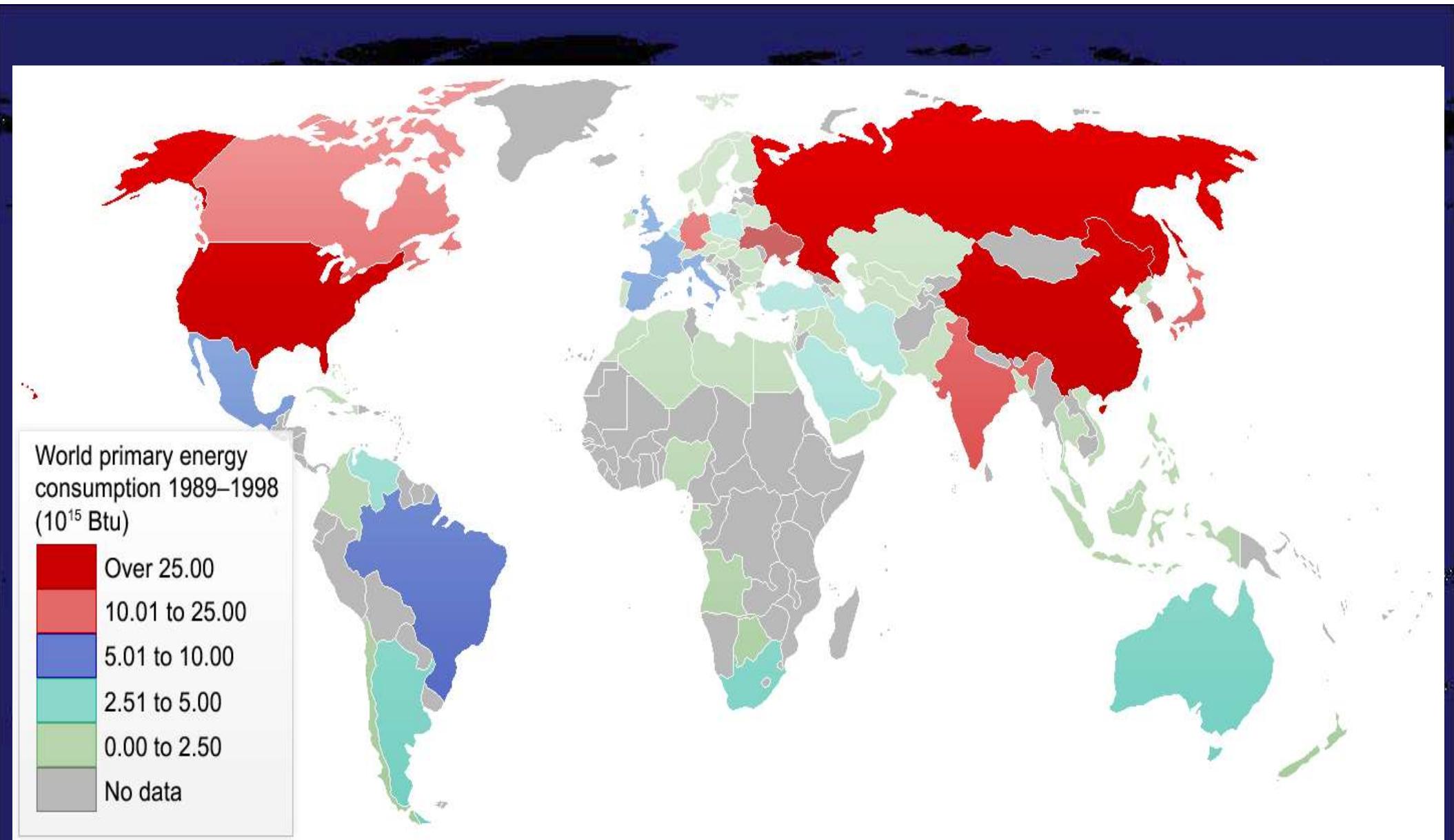
Introduction to climate change. Climate models and data. Statistical and dynamic downscaling of climatic data for use in regional scales. Exercise with hydrological simulation under climate change.

By

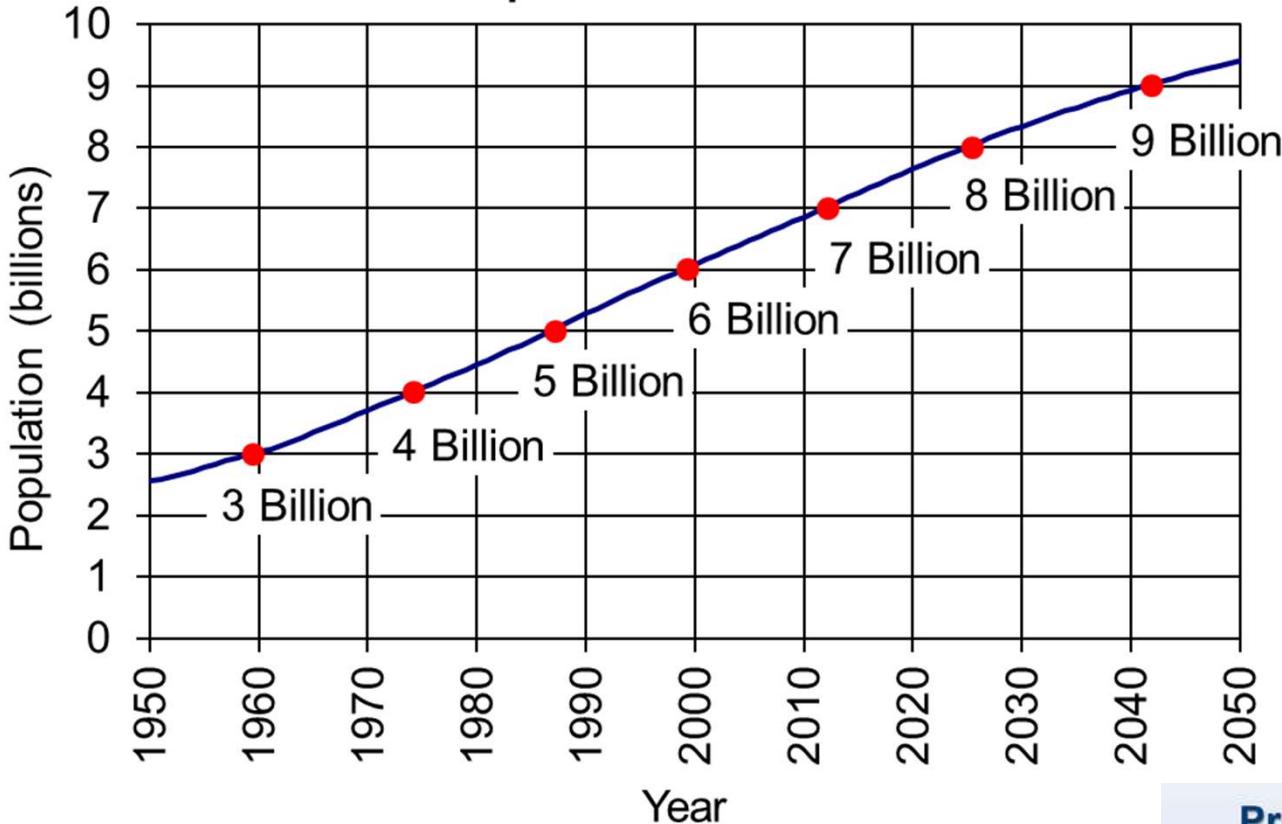
Dr. Charalampos (Haris) Skoulikaris
hskoulik@civil.auth.gr

Wednesday, 15/12/2021

Energy and development

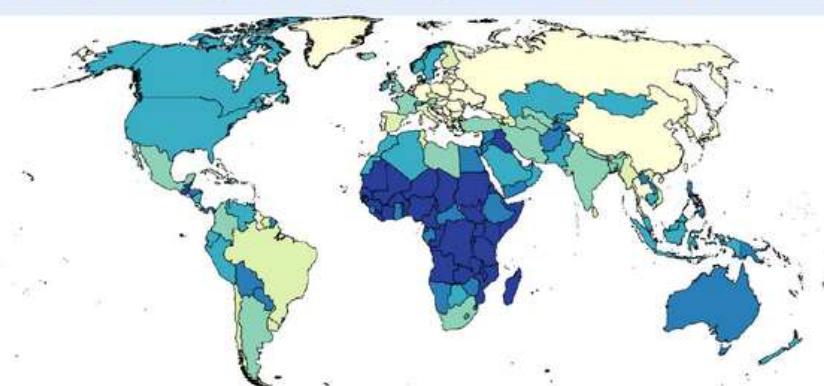


World Population: 1950-2050



Source: U.S. Census Bureau, International Data Base, July 2015 Update.

Projected population growth, 2010-2100

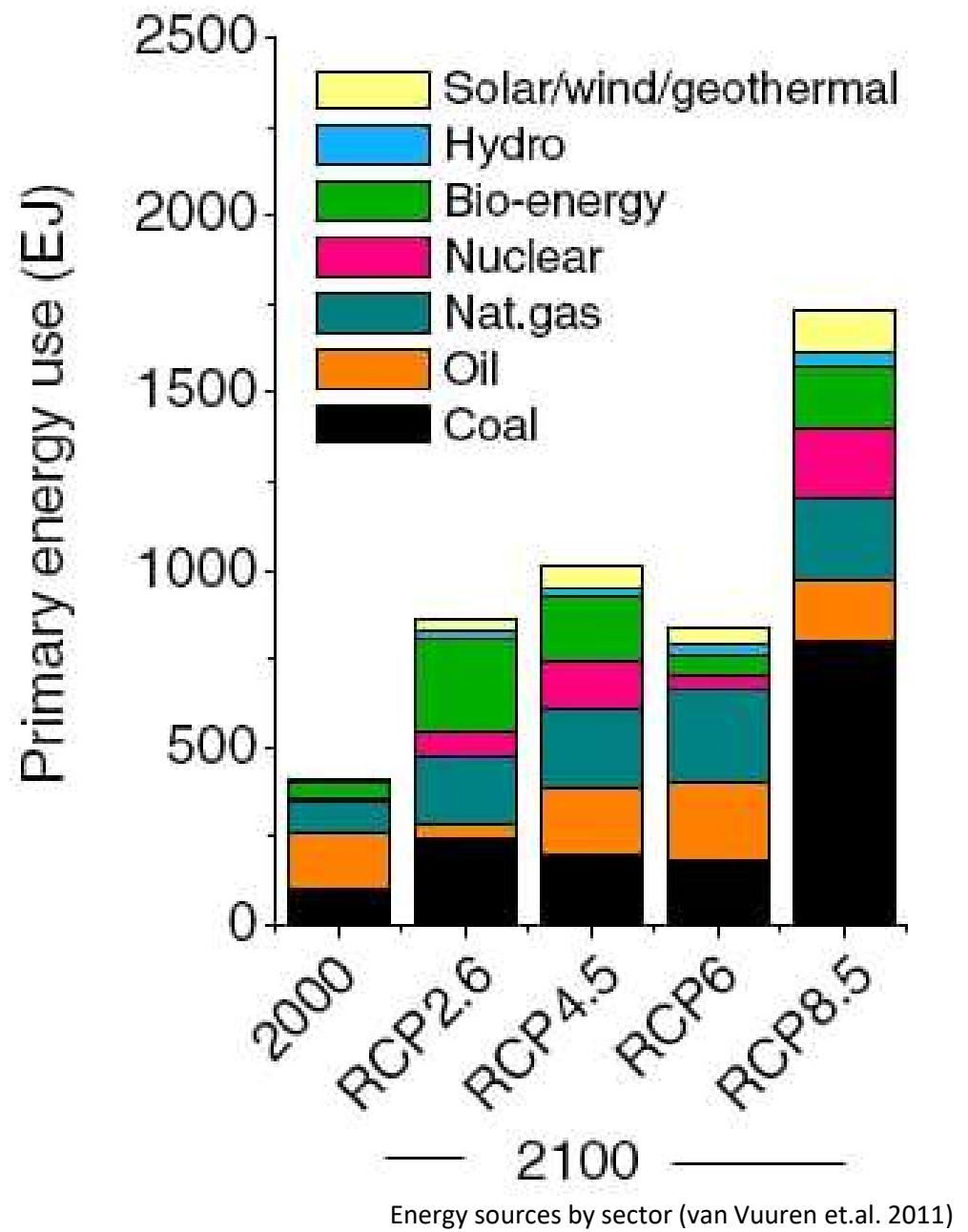
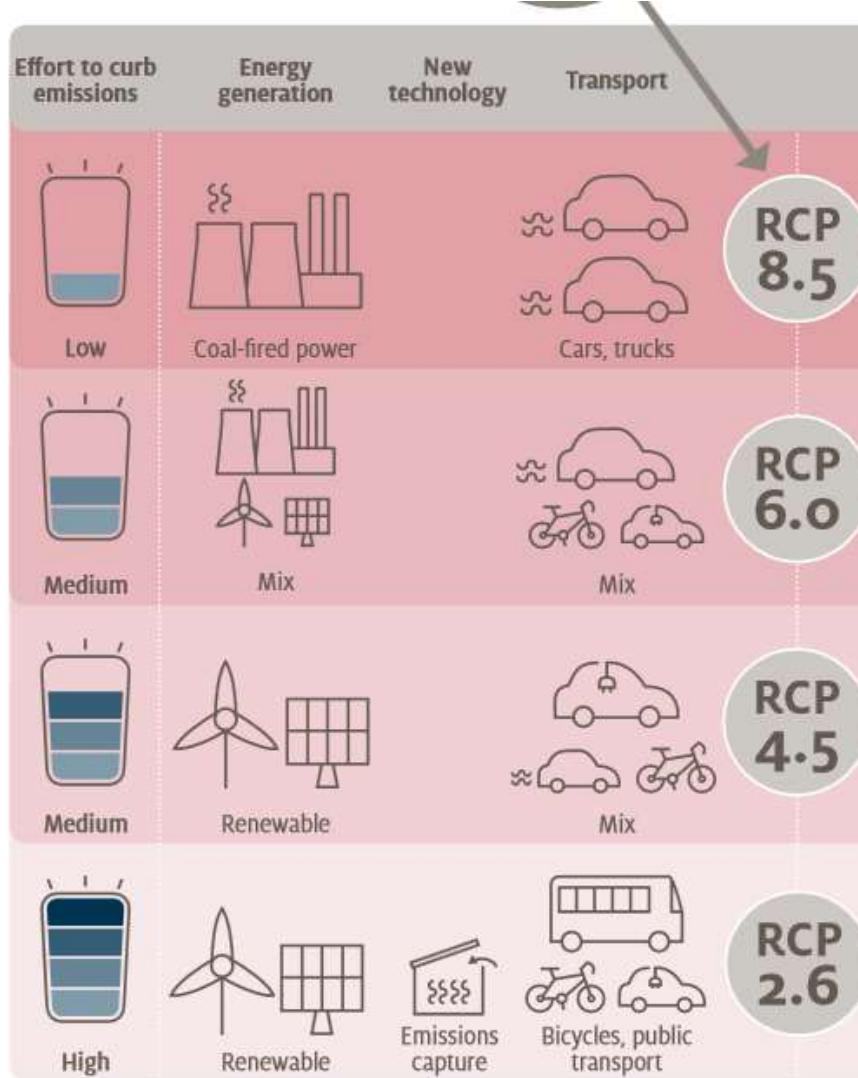


Percentage of population increase between 2010 and 2100

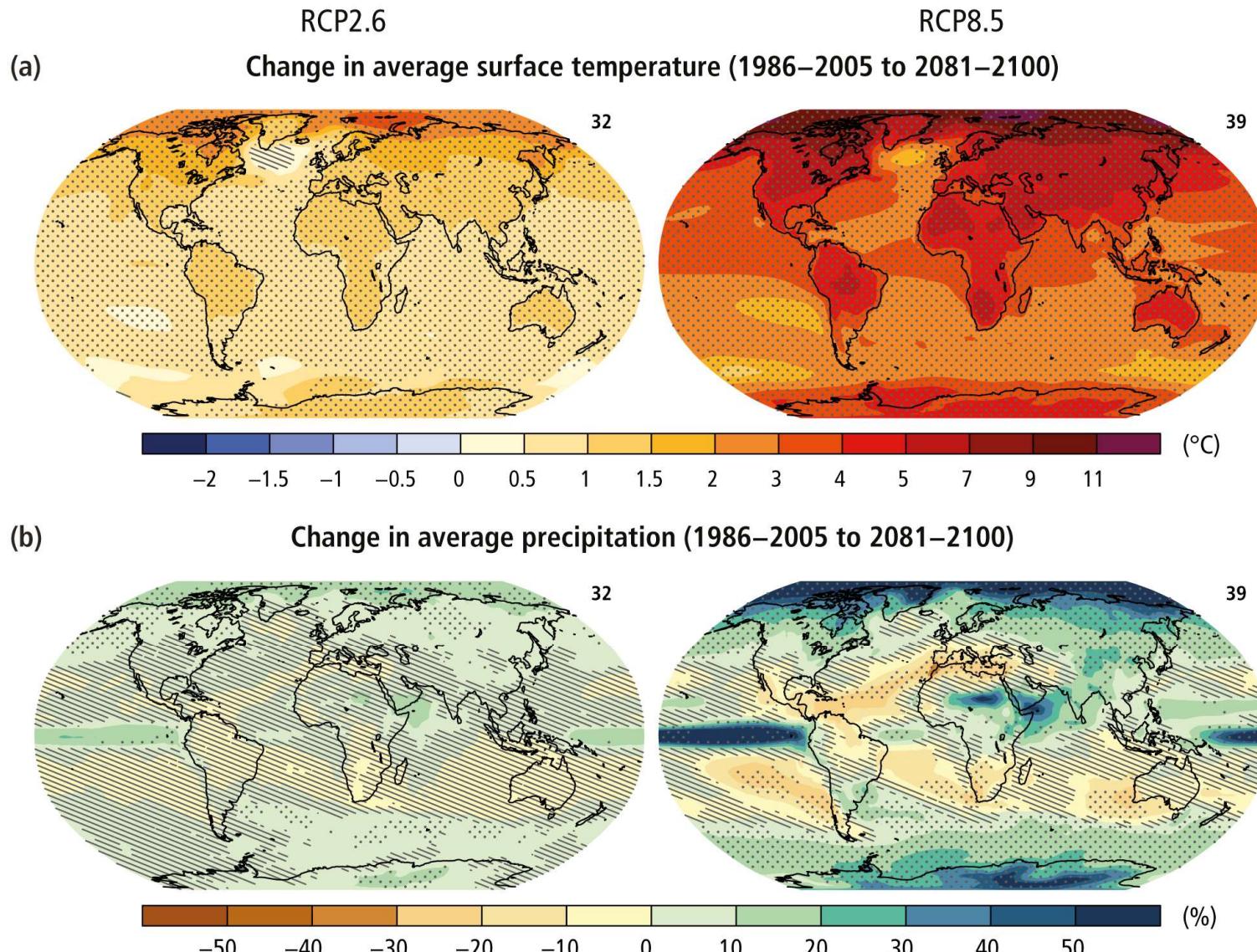
-60 - -20
-20 - 10
10 - 30
30 - 70
70 - 220
220 - 1200

Note: The boundaries on this map do not imply official endorsement or acceptance by the United Nations.

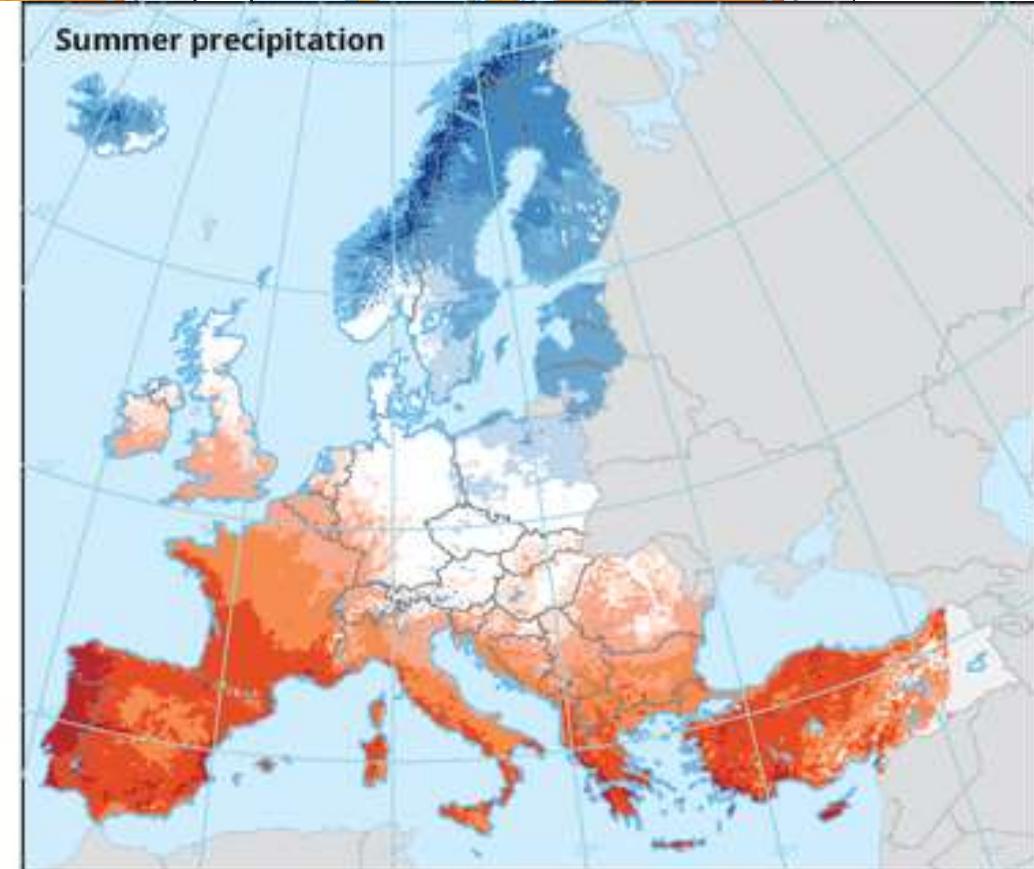
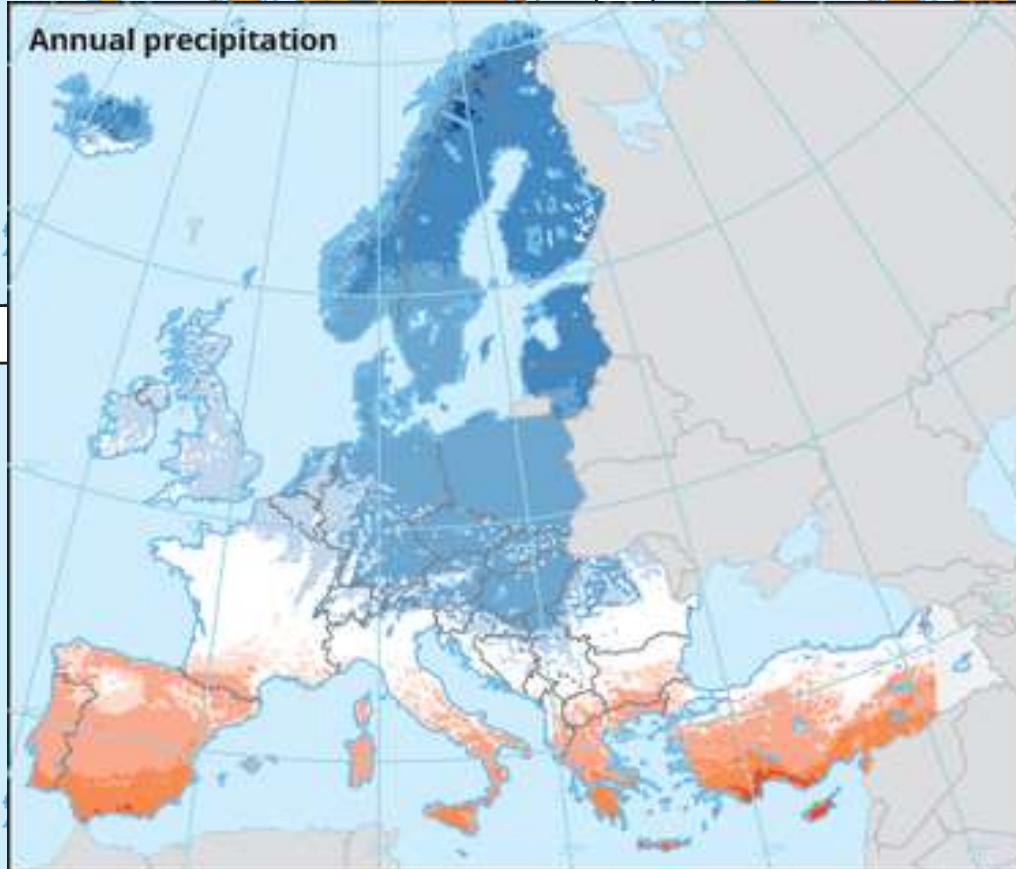
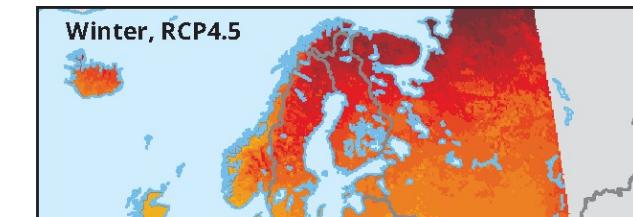
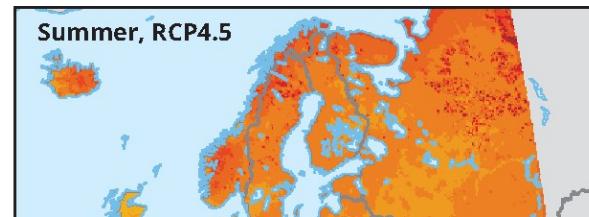
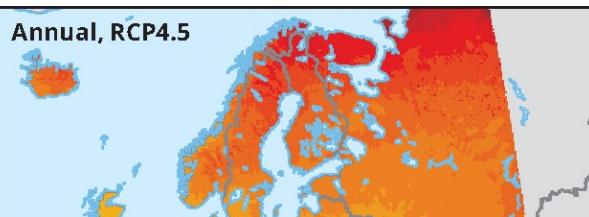
Source: United Nations, Department of Economic and Social Affairs, Population Division (2013). *World Population Prospects: The 2012 Revision*. New York. See www.unpopulation.org.



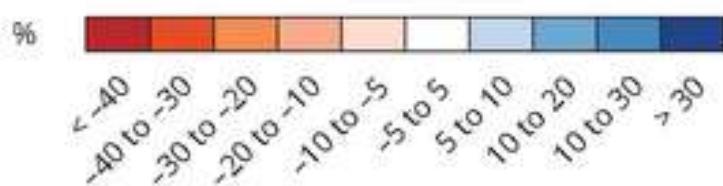
Changes in temperature and precipitation



Αλλαγής της μέσης επιφανειακής θερμοκρασίας **(a)** και αλλαγή της μέσης βροχόπτωσης **(b)** βάση της μέσης πρόγνωσης αριθμού μοντέλων για την περίοδο 2081–2100 σε σχέση με την 1986–2005 για τα σενάρια RCP2.6 (αριστερά) and RCP8.5 (δεξιά). Source: PCC AR5 WGI: Climate Change 2013: The Physical Science Basis, Technical Summary



Projected change in annual and summer precipitation, 2071-2100



No data
Outside scope

0 500 1000 1500 km

Reasons that cause Climate Change

Fossil Fuels



Industrial activities





Transportations-



Intense farming



Deforestation



Impacts of CC







FLOODS







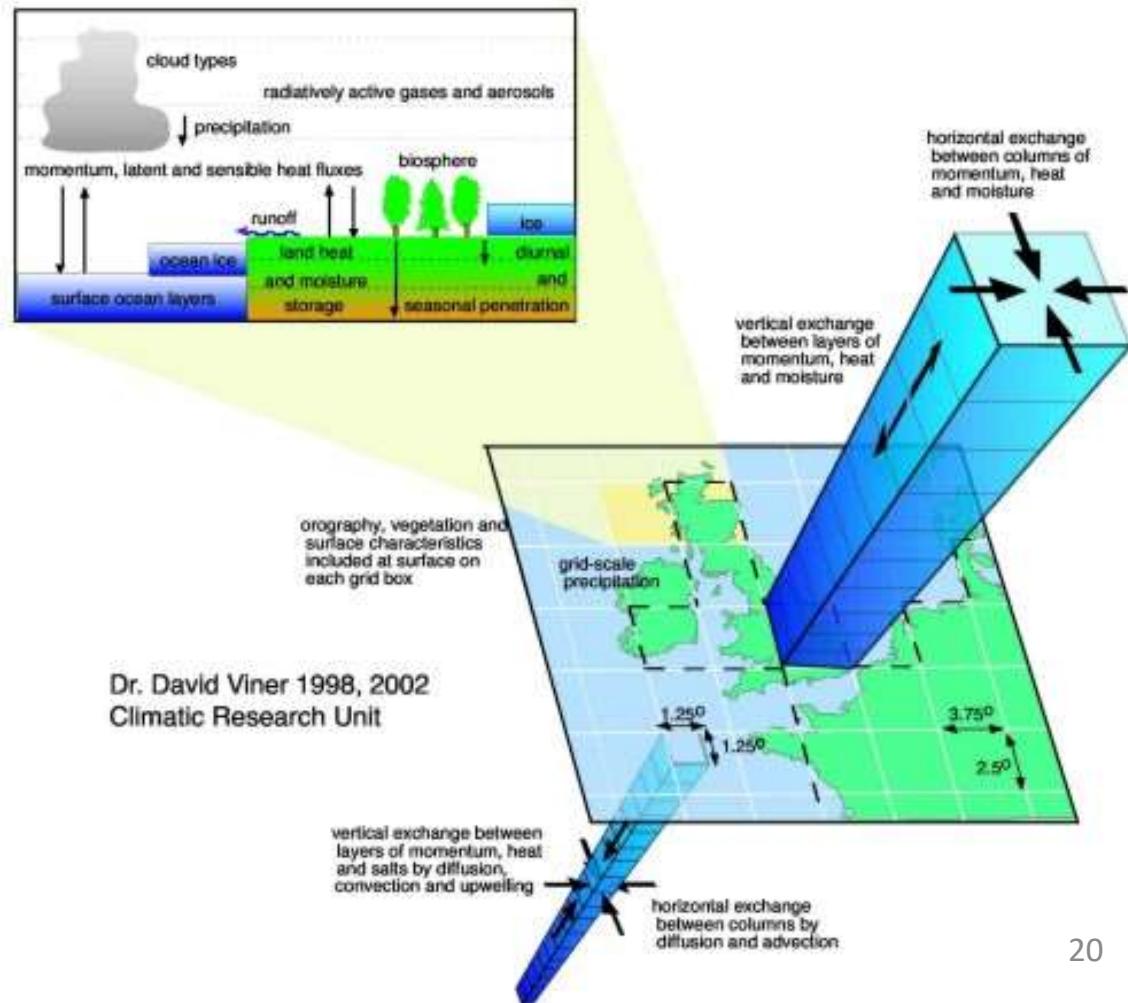
Population movements

General Circulation Models

Numerical models (General Circulation Models or GCMs), representing physical processes in the atmosphere, ocean, cryosphere and land surface, are the most advanced tools currently available for simulating the response of the global climate system to increasing greenhouse gas concentrations.

GCMs depict the climate using a 3 dimensional grid over the globe:

- a horizontal resolution of between 250 and 600 km,
- 10 to 20 vertical layers in the atmosphere and
- as many as 30 layers in the oceans.



Uncertainties in GCM-based simulations

1. Many physical processes, such as those related to clouds, also occur at smaller scales and cannot be properly modelled. Instead, their known properties must be averaged over the larger scale in a technique known as parameterization.

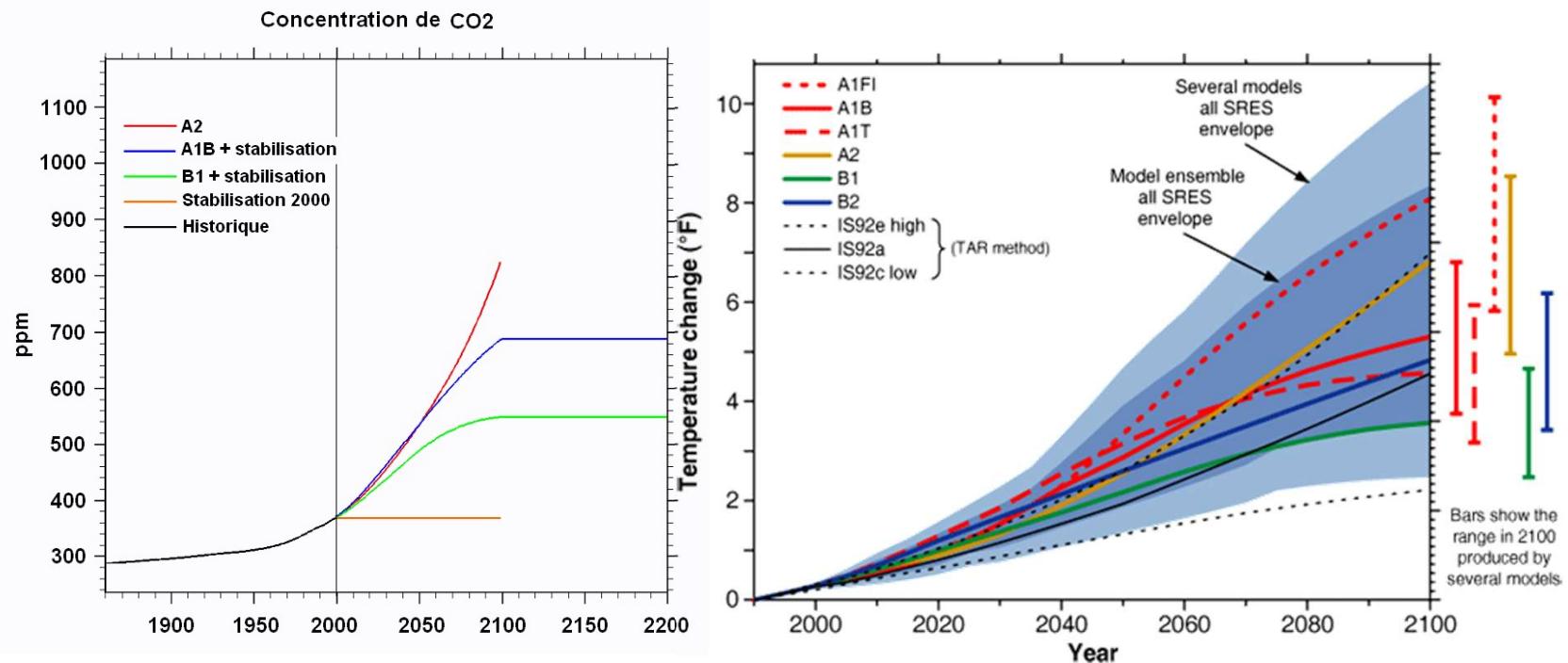
2. The simulation of various feedback mechanisms in models concerning, for example, water vapour and warming, clouds and radiation, ocean circulation and ice and snow albedo. For this reason, GCMs may simulate quite different responses to the same forcing, simply because of the way certain processes and feedbacks are modelled.

The climate change scenarios

The IPCC scenarios

In 1996 the Intergovernmental Panel on Climate Change (IPCC) created a report, the « **Special Report on Emission Scenarios** » (SRES). It describes the possible climate change scenarios: « stabilisation 2000 », B1, B2, A1, A2 and their variants.

- Bulgaria: Decrease of precipitations ?
- The IPCC scenarios



A1 : rapid economic growth
(groups: A1T; A1B; A1FI) **1.4 - 6.4 °C**

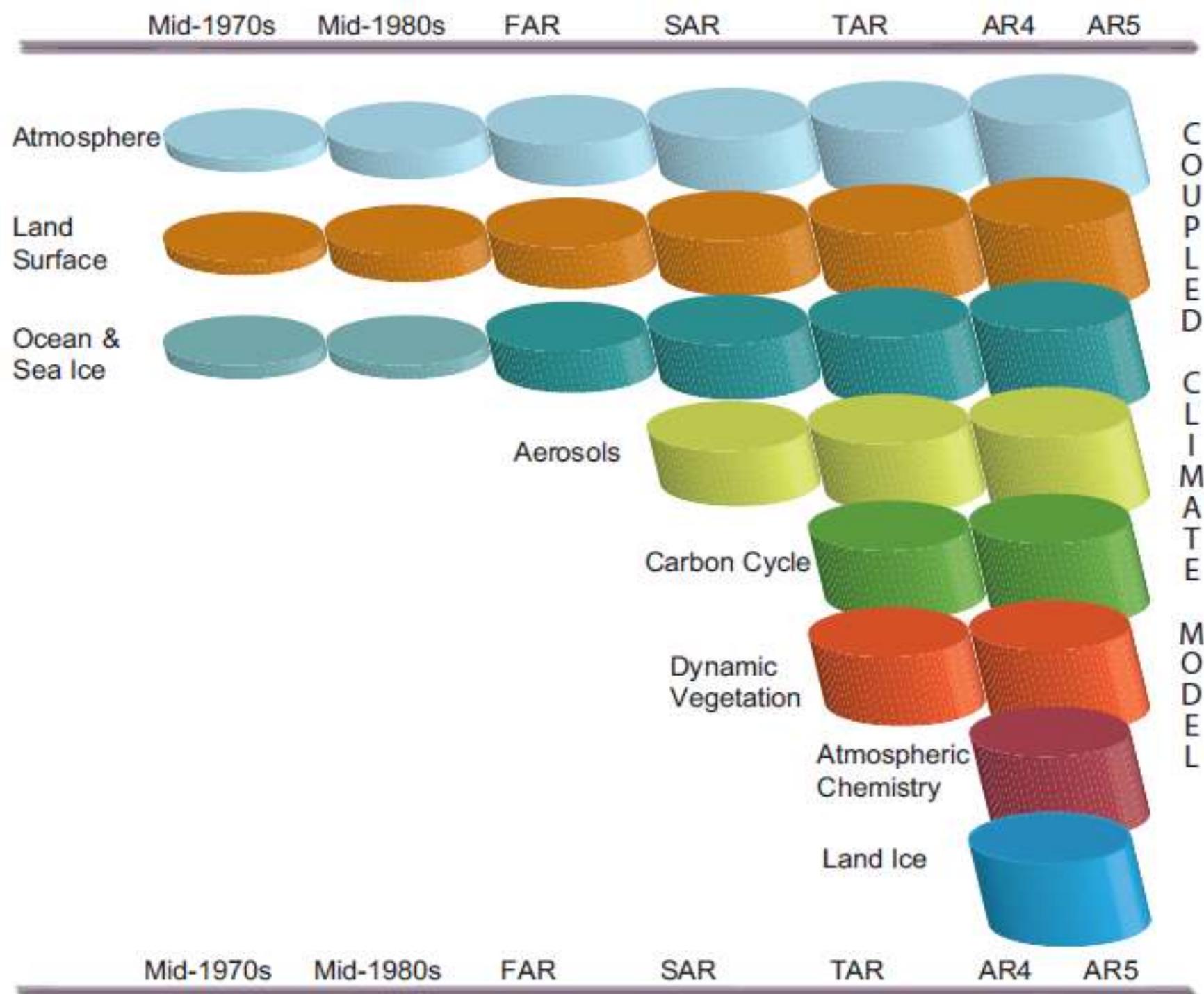
A2 : regionally oriented
economic development **2.0 - 5.4 °C**

B1: global environmental sustainability **1.1 - 2.9 °C**

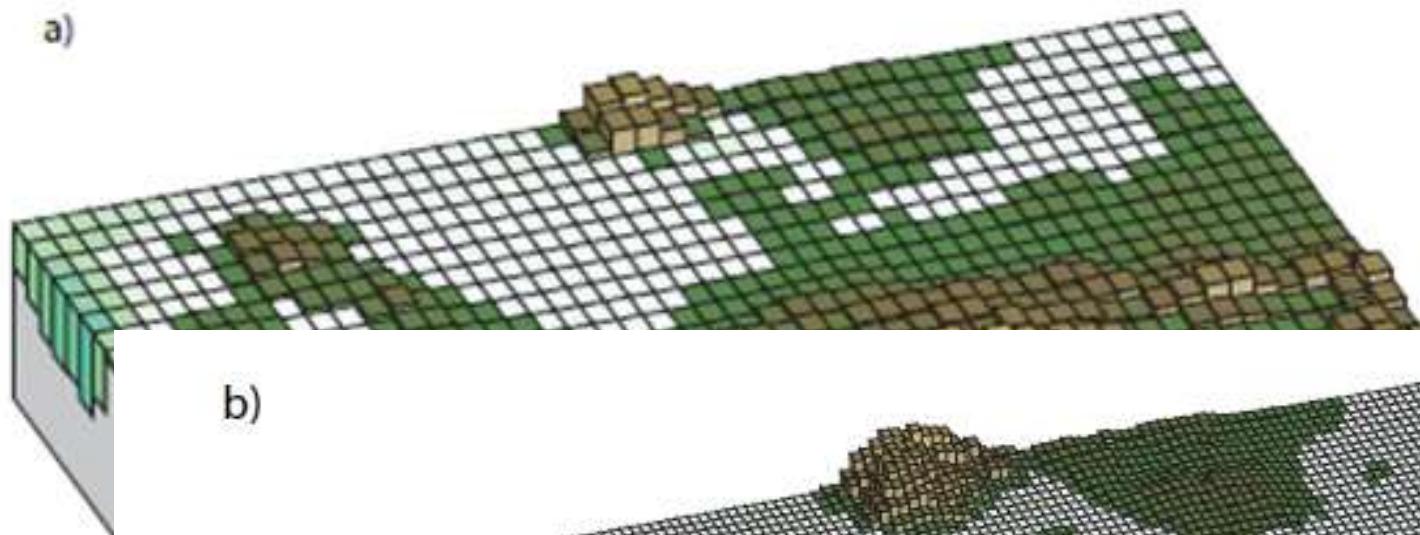
B2 : local environmental sustainability
1.4 - 3.8 °C

CC Assessment Reports

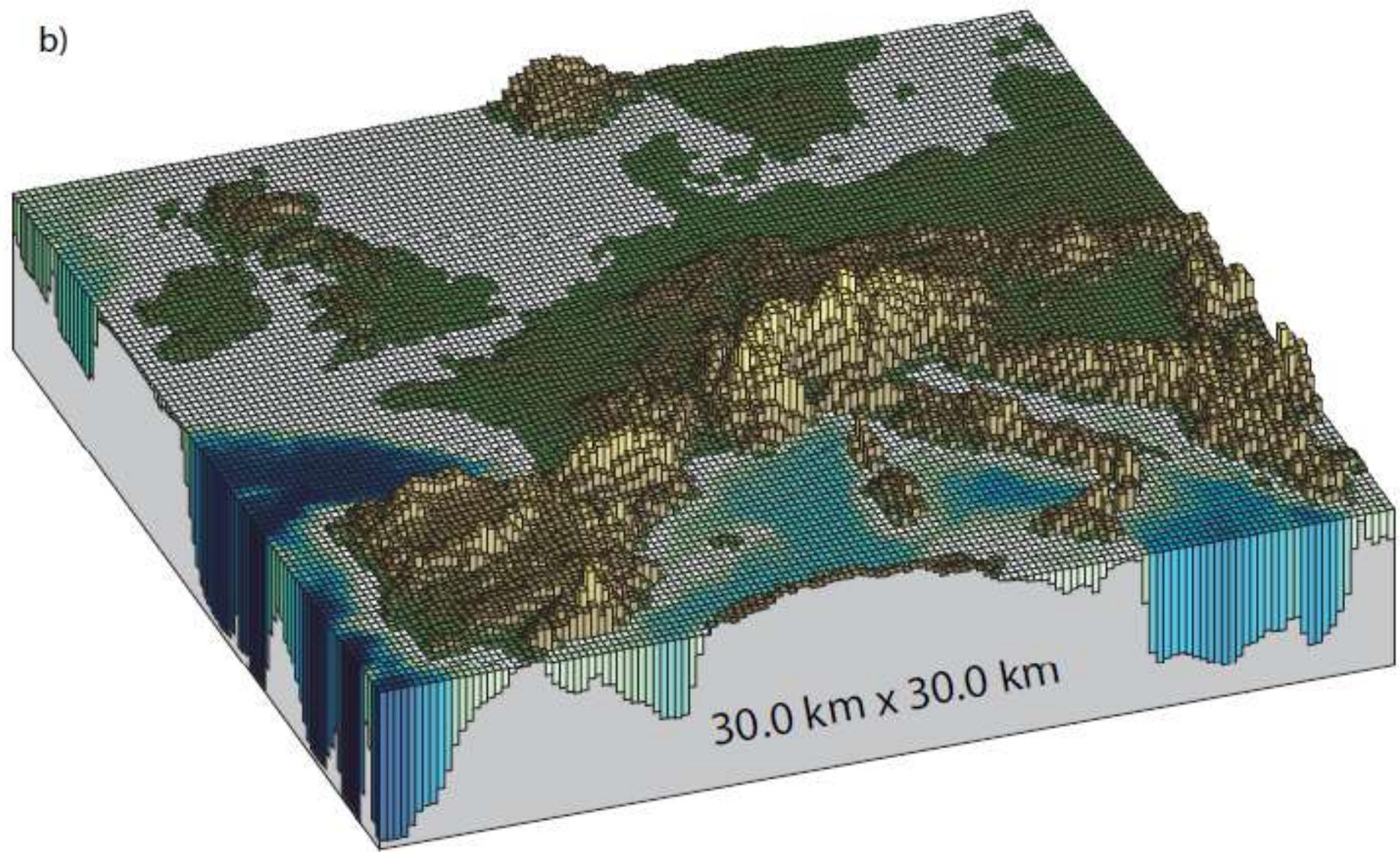
- *IPCC First Assessment Report 1990 (FAR)*
- *IPCC Second Assessment Report: Climate Change 1995 (SAR)*
- *IPCC Third Assessment Report: Climate Change 2001 (TAR)*
- *IPCC Fourth Assessment Report: Climate Change 2007 (AR4)*
- *IPCC Fifth Assessment Report: Climate Change 2013 (AR5)*



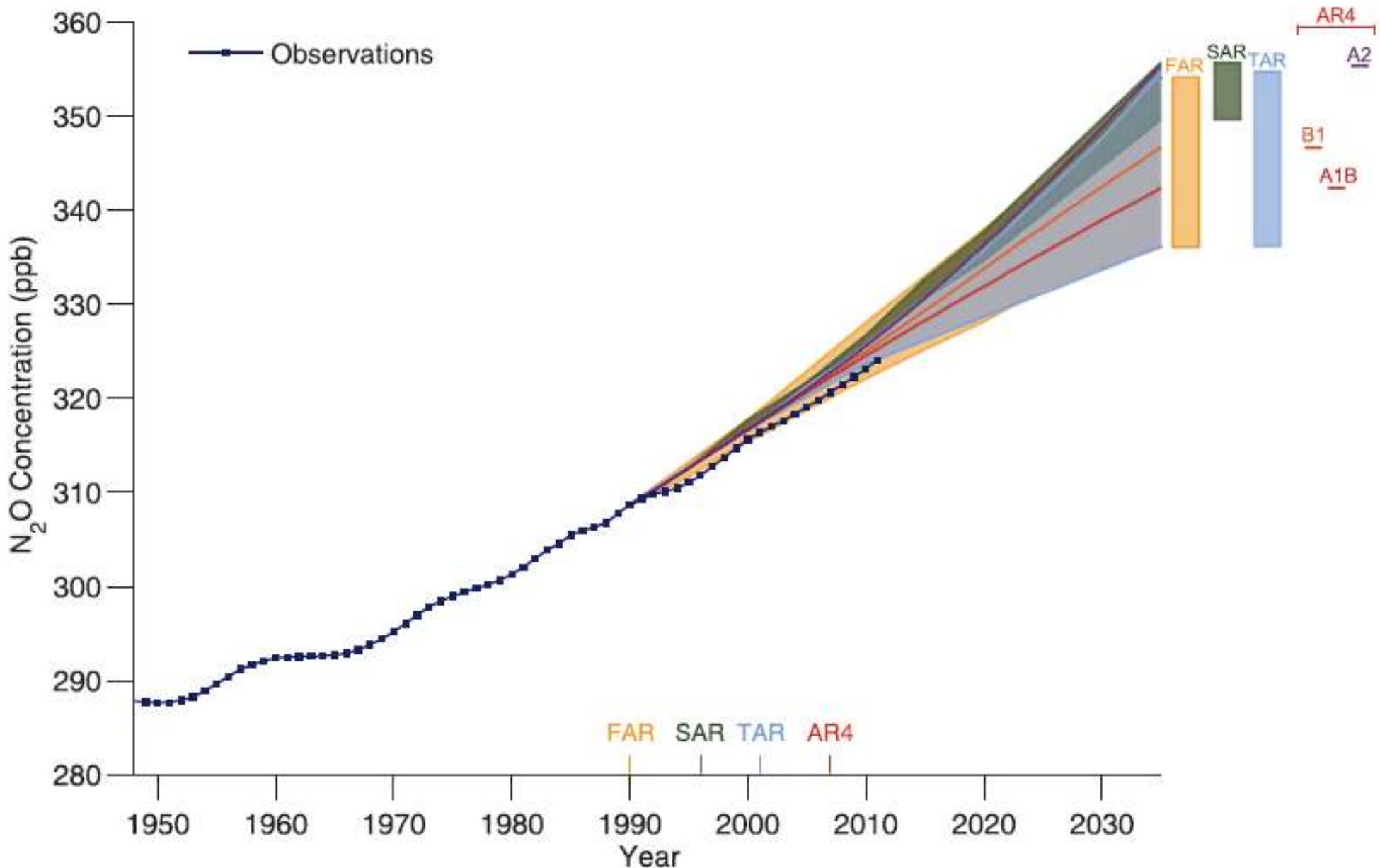
a)



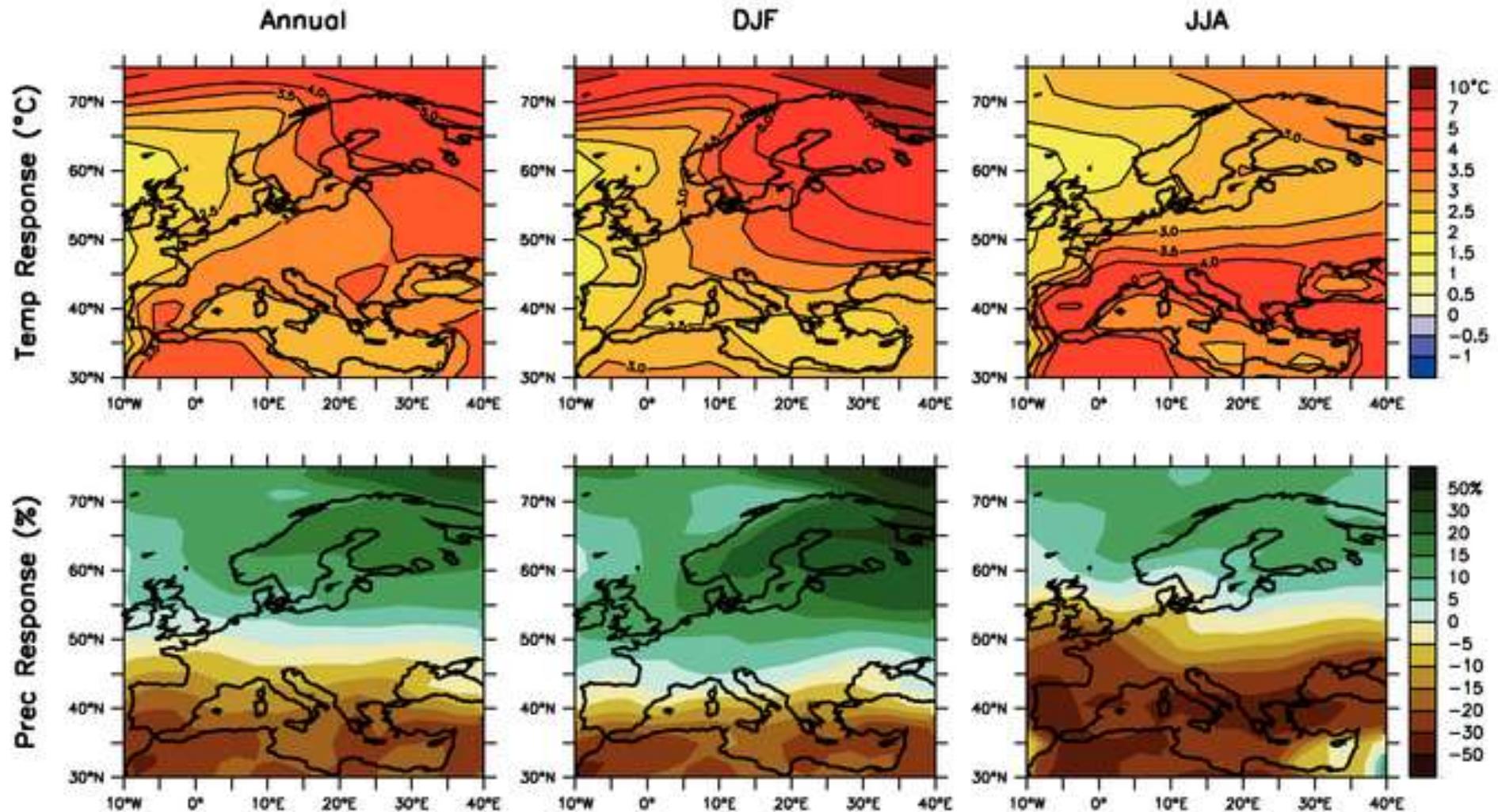
b)



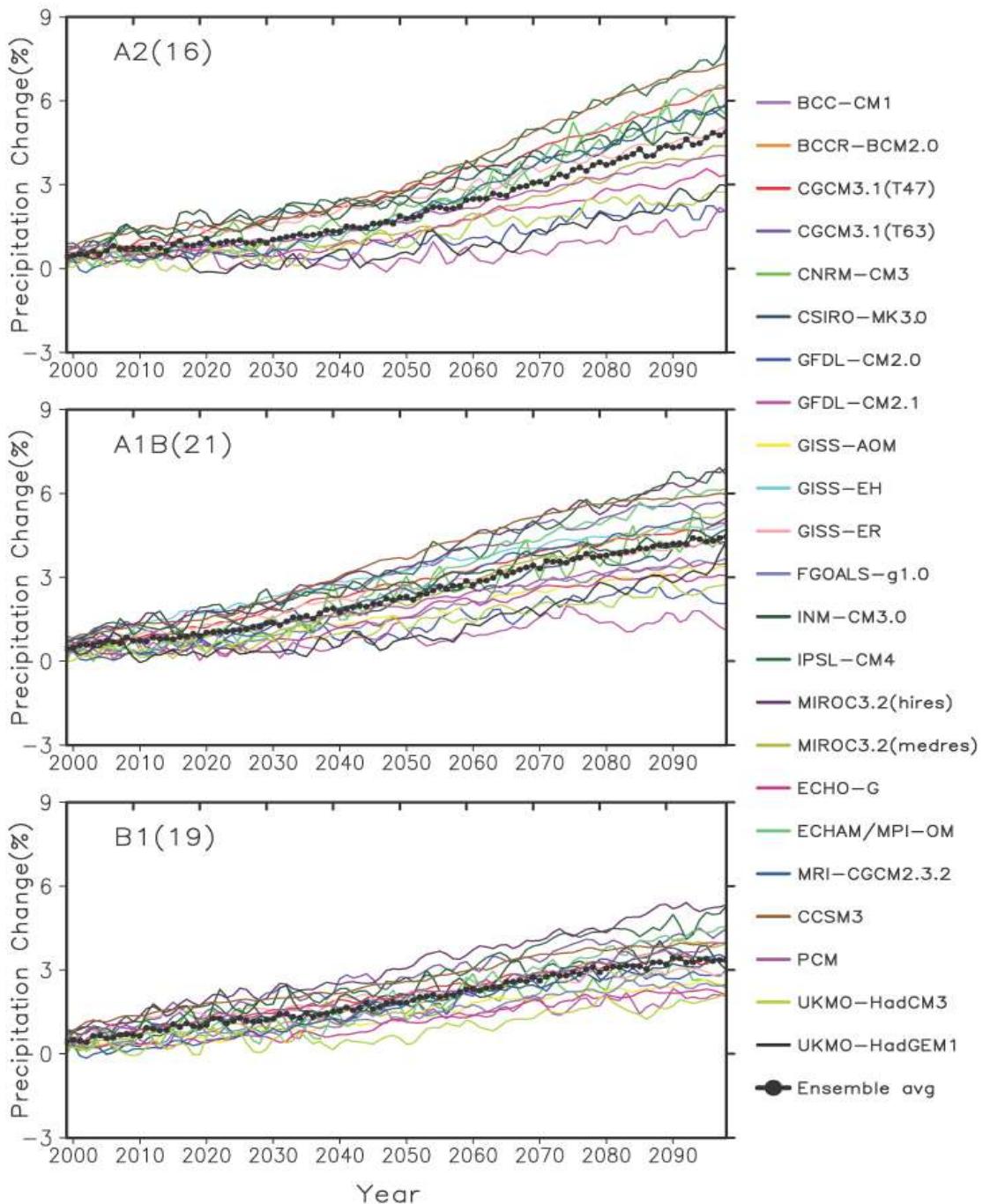
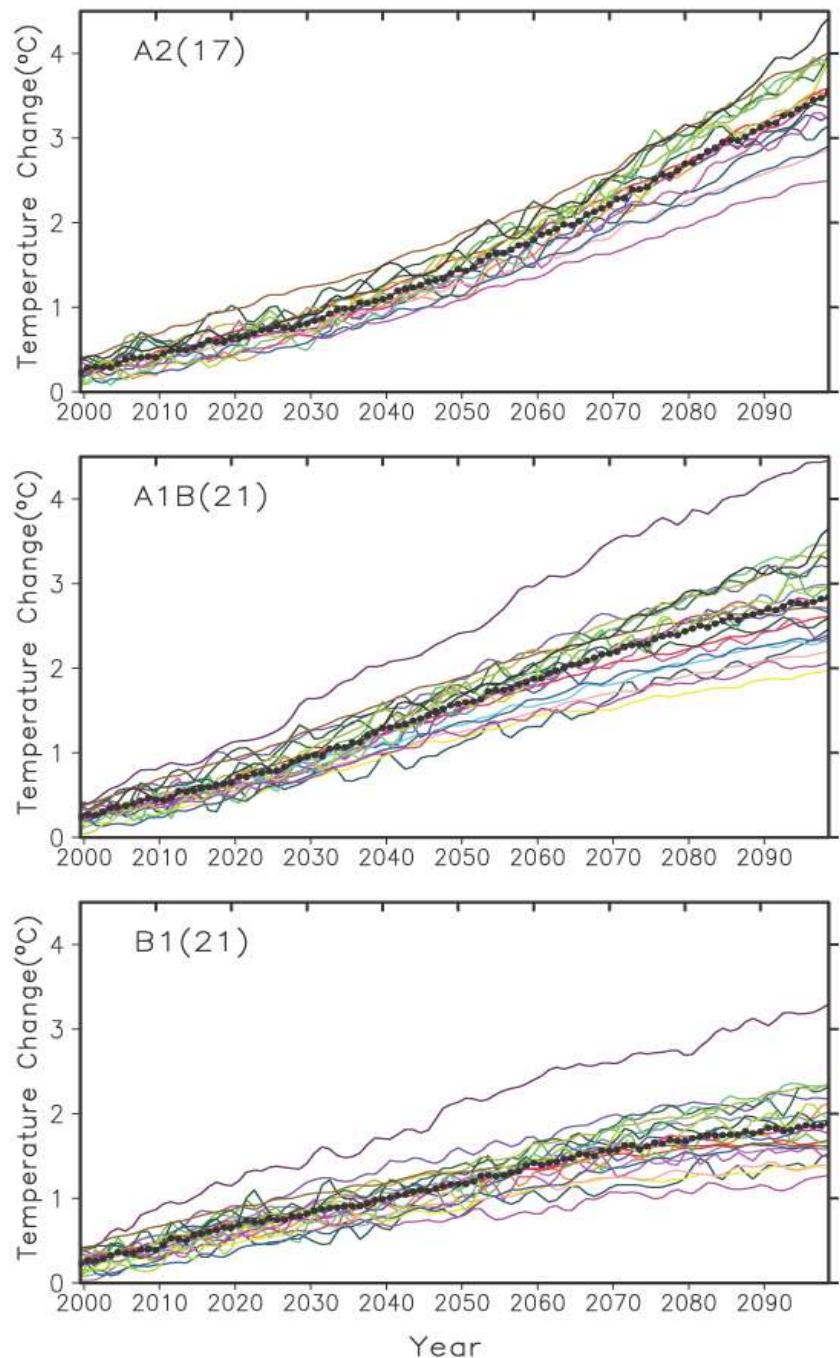
CO_2 , CH_4 , N_2O projections



Climate Change Projections

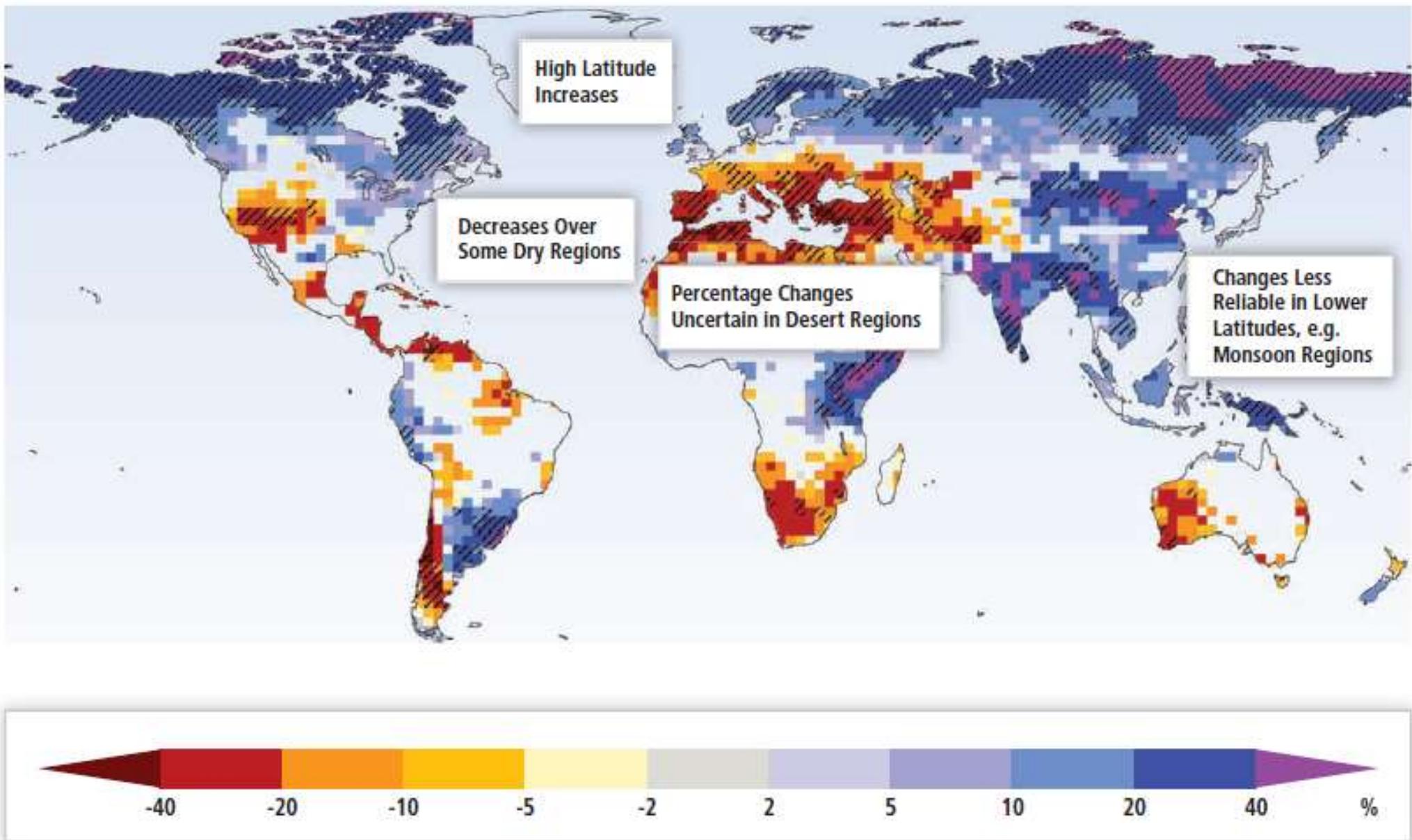


Temperature and precipitation changes over Europe from the MMD-A1B simulations, averaged over 21 models.



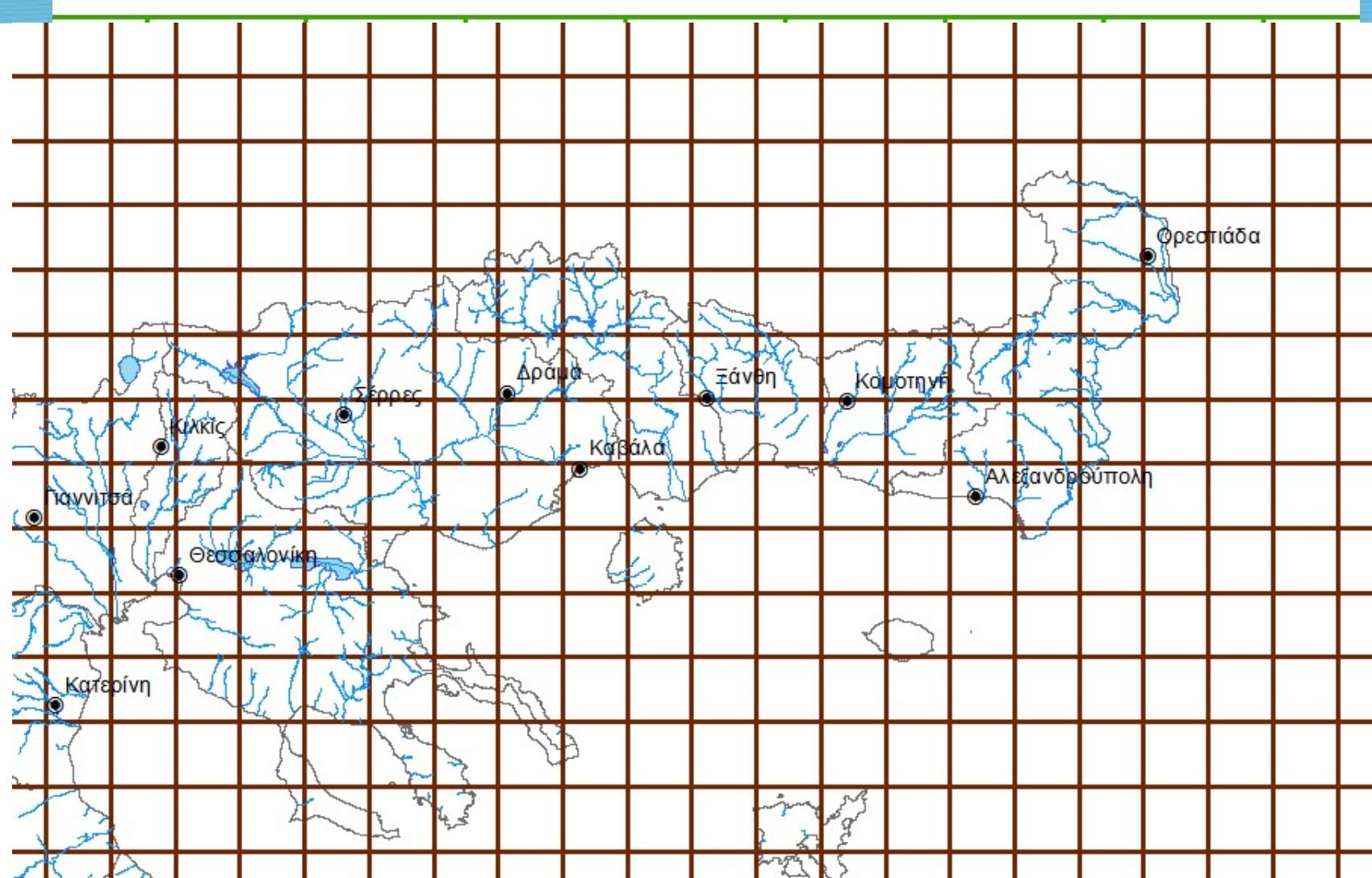
Source: IPCC Fourth Assessment Report: Climate Change 2007: WG I: The Physical science Basis

Changes in annual runoff



| Large-scale changes in annual runoff (water availability, in percent) for the period 2090 to 2099, relative to 1980 to 1999. Values represent the median of 12 climate model projections using the SRES A1B scenario. White areas are where less than 66% of the 12 models agree on the sign of change and hatched areas are where more than 90% of models agree on the sign of change. Source: IPCC (2007a).

Why do we need climate downscaling?



- **Red gird:** Grid of climate model with resolution of 200x200 km.
- **Green gird :** Grid of climate model with resolution of 50x50 km.
- **Brown grid:** Grid of climate model with resolution of 20x20 km.

What is Climate Downscaling?

Downscaling: obtaining subgrid-scale information from coarser resolution fields

Climate model downscaling “bridges the gap” between what is provided by global climate modelers and what is needed by engineers, scientists, decision-makers and impact assessors

**Two approaches: statistical
dynamical (regional climate models)**

Statistical downscaling

Statistical downscaling is based on the view that the regional climate is conditioned by two factors:

- the large scale climatic state, and
- the regional physiographic features (e.g. topography, land-sea distribution and land-use).

From this perspective, regional or local climate information is derived by first determining a statistical model which relates large scale climate variables (or “**predictors**”) to regional and local variables (or “**predictands**”). Then the large scale output of a GCM simulation is fed into this statistical model to estimate the corresponding local and regional climate characteristics.

Statistical downscaling involves developing quantitative relationships between large scale atmospheric variables (predictors) and local surface variables (predictands). The most common form has the predictand as a function of the predictor(s), but other types of relationships have been used (e.g. between predictors and the statistical distribution parameters of the predictand, or the frequencies of extremes of the predictands)

SD methods classification & description

- **Weather classification** methods group days into finite number of discrete weather types or “states” according to their synoptic similarity. The predictand is then assigned to the prevailing weather state, and replicated under changed climate conditions by resampling or regression functions.
- **Regression models** are a conceptually simple means of representing linear or linear relationships between predictands and the large scale atmospheric forcing. (problem with daily precipitation downscaling because of the relatively low relationships)
- **Weather generators** are models that replicate the statistical attributes of a local climate variable (such as mean and variance) but not observed sequences of events.

+ & - of SD methods

Method	Strengths	Weaknesses
Weather typing (e.g. analogue method, hybrid approaches, fuzzy classification, self organizing maps, Monte Carlo methods).	<ul style="list-style-type: none"> • Yields physically interpretable linkages to surface climate • Versatile (e.g., can be applied to surface climate, air quality, flooding, erosion, etc.) • Compositing for analysis of extreme events 	<ul style="list-style-type: none"> • Requires additional task of weather classification • Circulation-based schemes can be insensitive to future climate forcing • May not capture intra-type variations in surface climate
Weather generators (e.g. Markov chains, stochastic models, spell length methods, storm arrival times, mixture modelling).	<ul style="list-style-type: none"> • Production of large ensembles for uncertainty analysis or long simulations for extremes • Spatial interpolation of model parameters using landscape • Can generate sub-daily information 	<ul style="list-style-type: none"> • Arbitrary adjustment of parameters for future climate • Unanticipated effects to secondary variables of changing precipitation parameters
Regression methods (e.g. linear regression, neural networks, canonical correlation analysis, kriging).	<ul style="list-style-type: none"> • Relatively straightforward to apply • Employs full range of available predictor variables • ‘Off-the-shelf’ solutions and software available 	<ul style="list-style-type: none"> • Poor representation of observed variance • May assume linearity and/or normality of data • Poor representation of extreme events

References and documentation

Guidelines for Use of Climate Scenarios Developed from Statistical Downscaling Methods

RL Wilby^{1,2}, SP Charles³, E Zorita⁴, B Timbal⁵, P Whetton⁶, LO Mearns⁷

¹Environment Agency of England and Wales, UK

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³CSIRO Land and Water, Australia

⁴GKSS, Germany

⁵Bureau of Meteorology, Australia

⁶CSIRO Atmospheric Research, Australia

⁷National Center for Atmospheric Research, USA

August 2004

The screenshot shows the UNFCCC website with a search bar and social media links at the top. Below is a navigation menu with Home, CDM, JI, CCiNet, TT:Clear, and Your location: Home. A blue box on the left says 'NEWSROOM Get News on the Latest Climate Action'. To the right, a table titled 'Compendium on methods and tools to evaluate impacts of, and vulnerability and adaptation to, climate change' has a row for 'Statistical DownScaling Model (SDSM)'.

Statistical DownScaling Model (SDSM)	
Description	SDSM is a user-friendly software package designed to implement statistical downscaling methods to produce high-resolution monthly climate information from coarse-resolution climate model (GCM) simulations. The software also uses weather generator methods to produce multiple realizations (ensembles) of synthetic daily weather sequences.
Appropriate Use	SDSM can be used whenever impact assessments require small-scale climate scenarios, provided quality observational data and daily GCM outputs for large-scale climate variables are available.
Scenarios	All scenarios

A COMPARISON OF STATISTICAL DOWNSCALING AND CLIMATE CHANGE FACTOR METHODS: IMPACTS ON LOW FLOWS IN THE RIVER THAMES, UNITED KINGDOM

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³Department of Geography, Kings College, London, United Kingdom

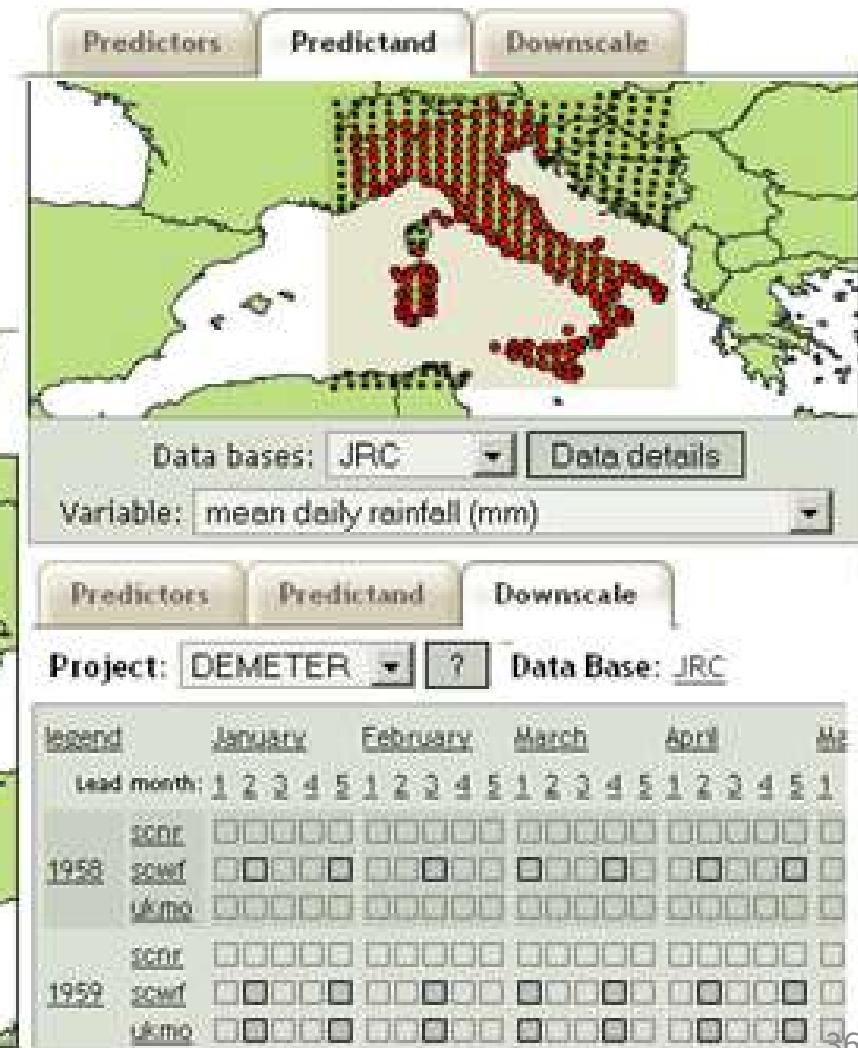
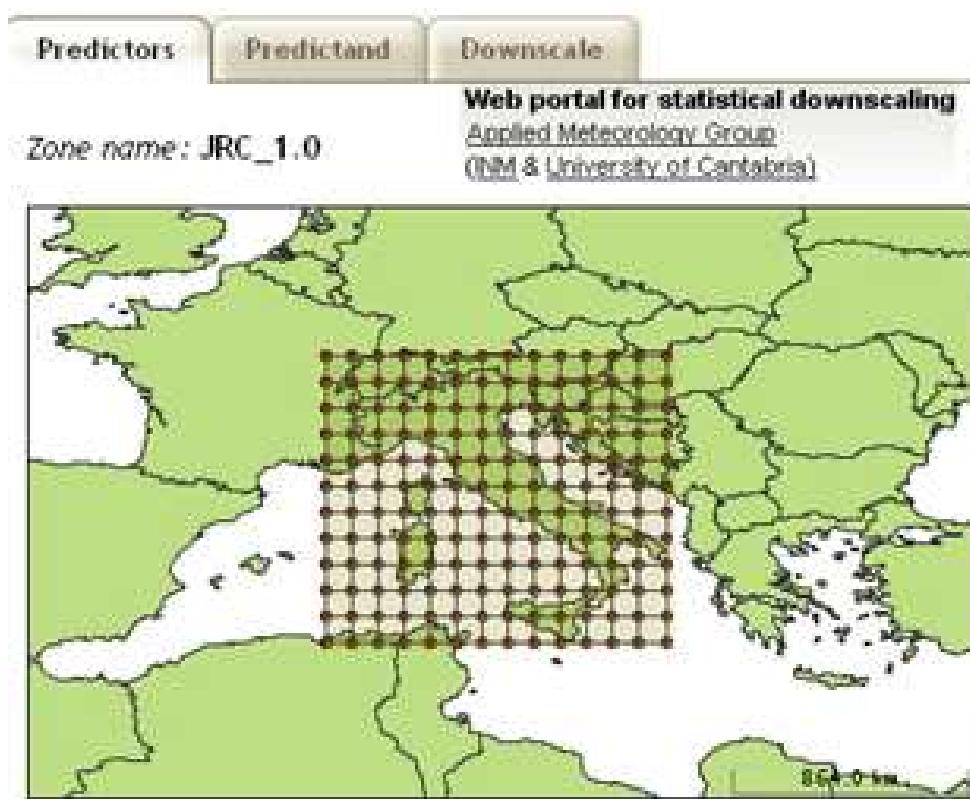
Wilby, R. L., Wigley, T. M. L., Conway, D., Jones, P. D., Hewitson, B. C., Main, J., & Wilks, D. S. (1998). Statistical downscaling of general circulation model output: a comparison of methods. *Water resources research*, 34(11), 2995-3008.

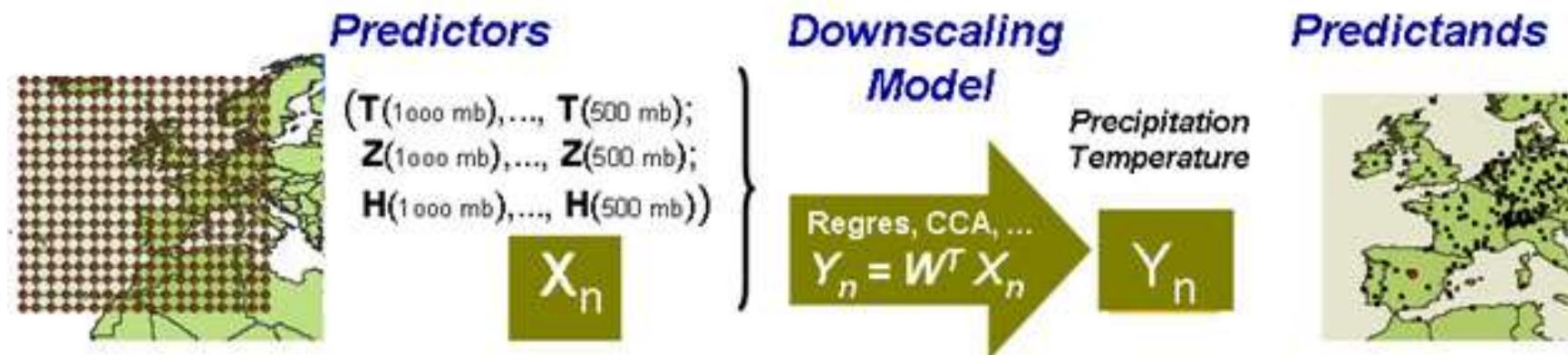
Wilby, R. L., Hay, L. E., Gutowski, W. J., Arritt, R. W., Takle, E. S., Pan, Z., ... & Clark, M. P. (2000). Hydrological responses to dynamically and statistically downscaled climate model output. *Geophysical Research Letters*, 27(8), 1199-1202.

On line models for SD

Three steps are necessary to obtain high resolution forecasts in a region of interest:

1. Selecting the predictors,
2. Selecting the stations and variable,
3. Running the desired downscaling jobs





The **main features** are:

- **PREDICTANDS:** JRC (0.5°) grid over Europe. ECA network (1000 stations) over Europe.
- **PREDICTORS:**
Climate Change predictions: ENSEMBLES (ECHAM5-MPI Germany), PCMDI-IPCC (CSIRO-MK3 Australia; CGCM3 Canada), local providers (CNRM-CM3).
- **DOWNSCALING:** analogs and weather typing, regression and neural networks

Dynamic downscaling

Nesting a regional climate model into an existing GCM is known as a dynamic method to downscale data.

To do this, a specific location is defined and certain driving factors from the GCM are applied to the regional climate model.

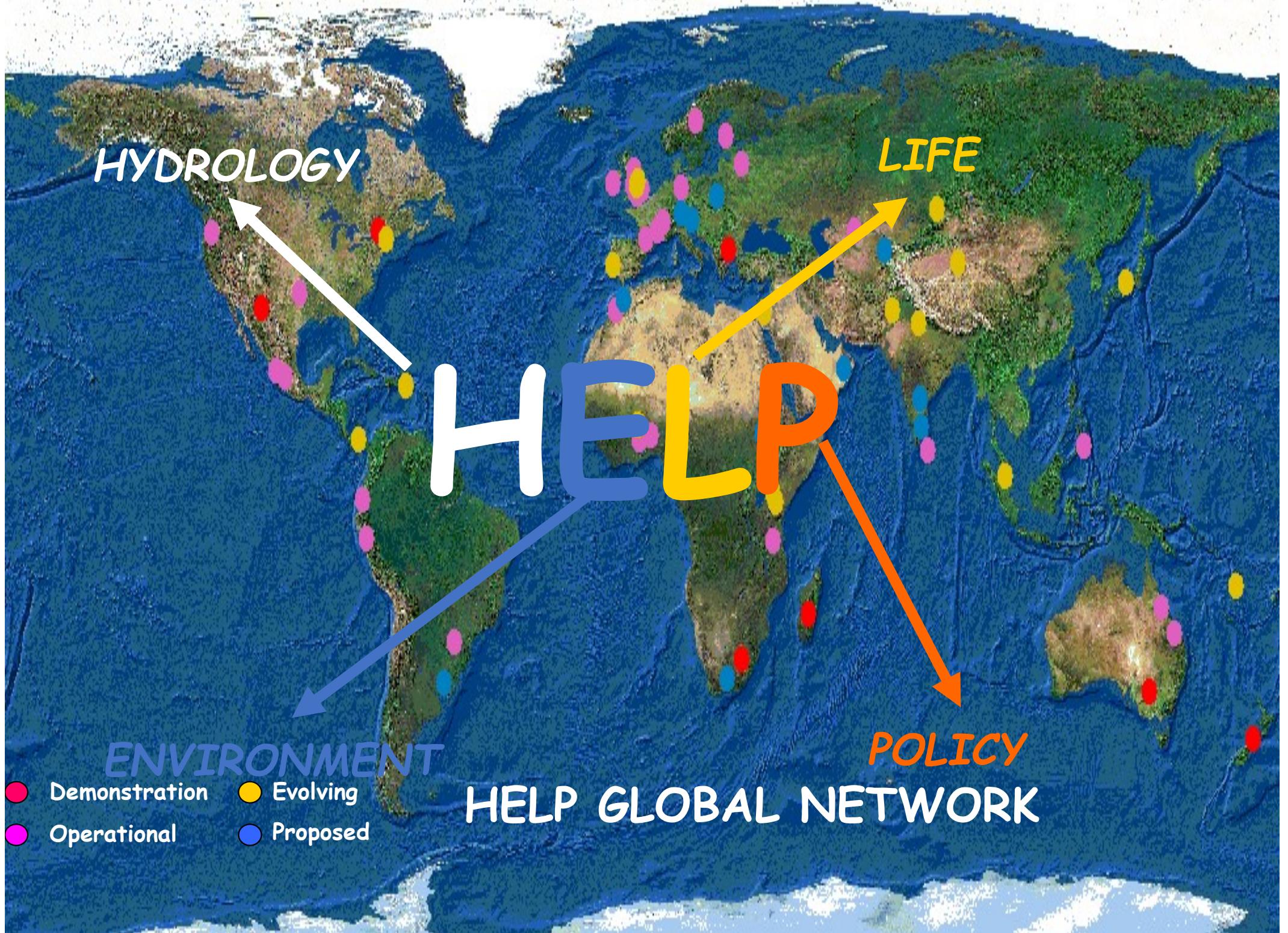
A regional climate model is a dynamic model, like a GCM, but it can be thought of as being composed of three layers:

- One layer is largely driven by the GCM,
- another layer builds on some locally specific data, and
- the third layer uses its own physics based equations to resolve the model based on data from the other two.

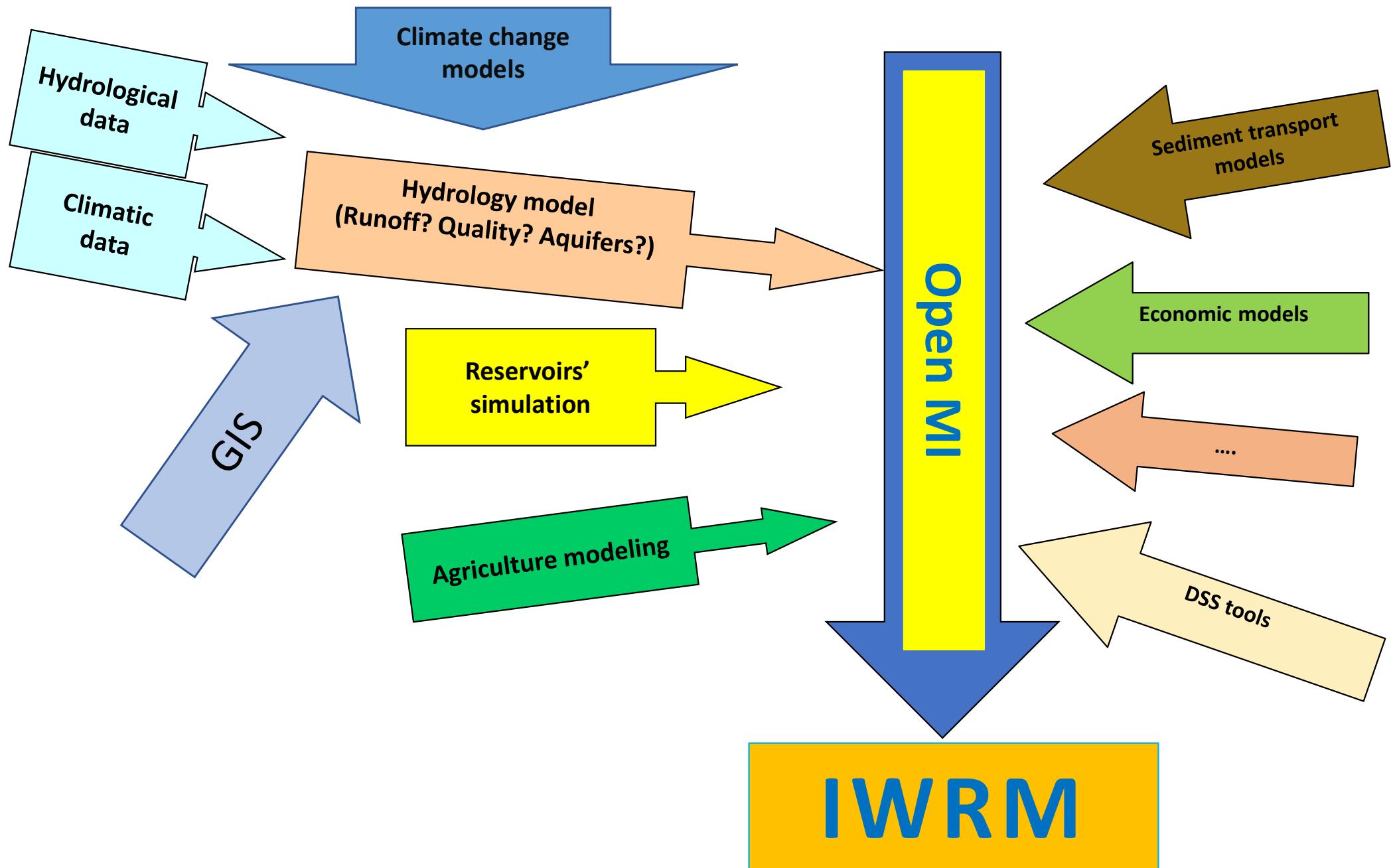
Climate change potential impacts to HPP

The resource potential for hydropower is currently based on historical data for the present climatic conditions. With a changing climate, this resource potential could change due to:

- **Changes in river flow** (runoff) related to changes in local climate, particularly in precipitation and temperature in the catchment area. This may lead to changes in runoff volume, variability of flow and seasonality of the flow (e.g., by changing from spring/summer high flow to more winter flow), directly affecting the resource potential for hydropower generation.
- **Changes in extreme events** (floods and droughts) may increase the cost and risk for the hydropower projects.
- **Changes in sediment loads** due to changing hydrology and extreme events. More sediment could increase turbine abrasions and decrease efficiency. Increased sediment load could also fill up reservoirs faster and decrease the live storage, reducing the degree of regulation and decreasing storage services.

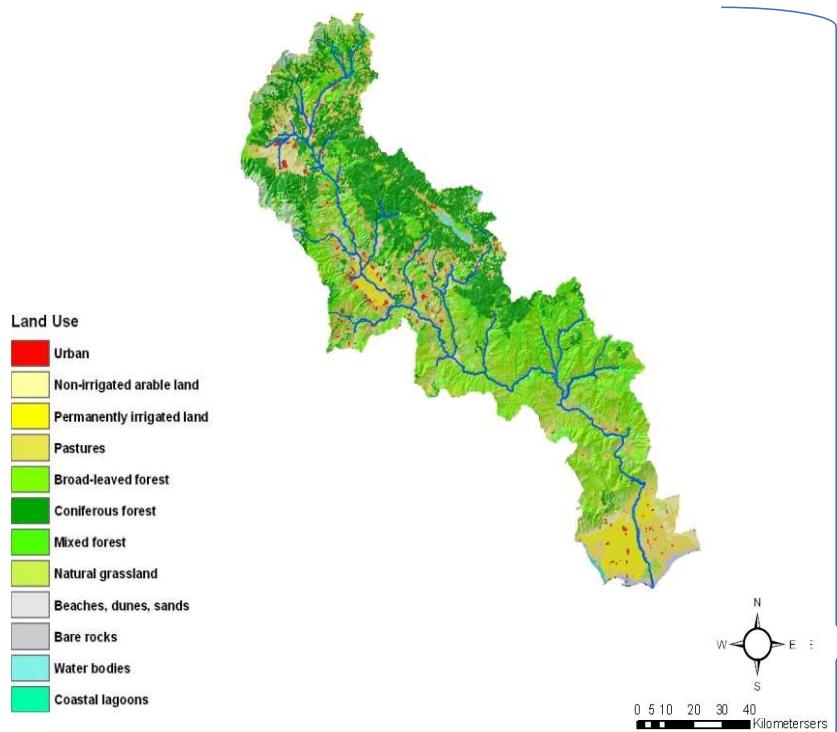


Modeling Coupling for the Integrated Water Resources Management

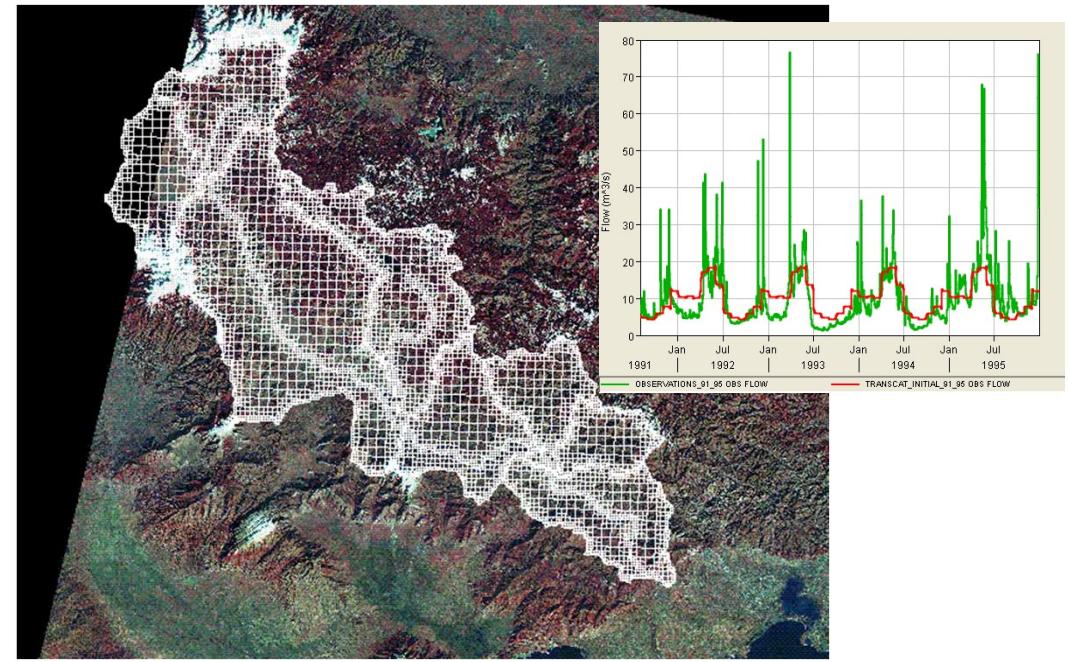
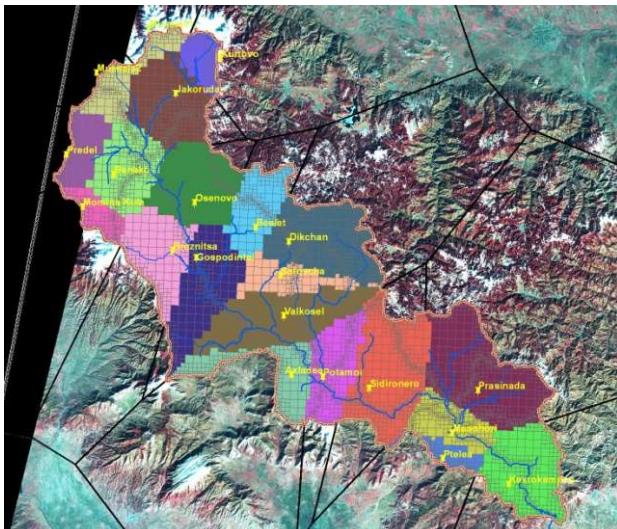


Geographic Information Systems - GIS

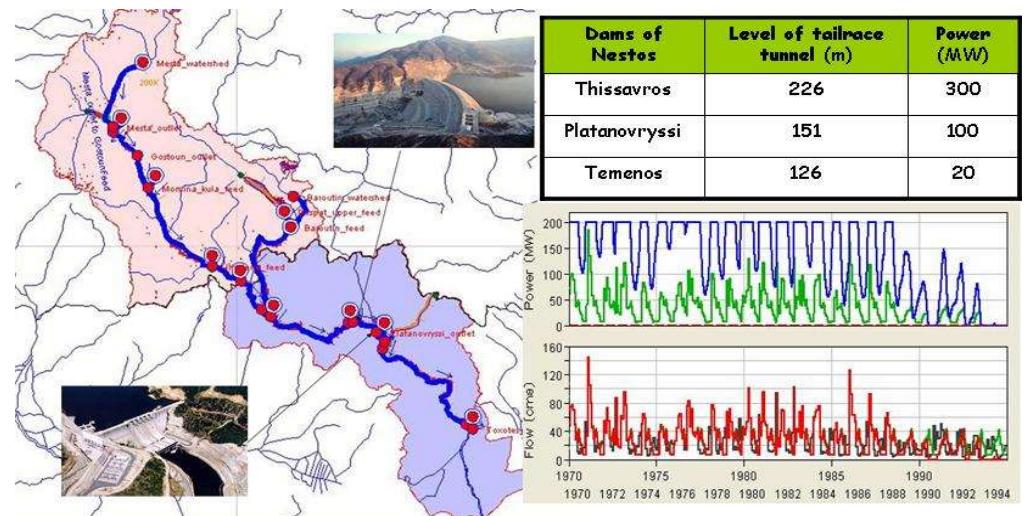
Hydrologic modelling



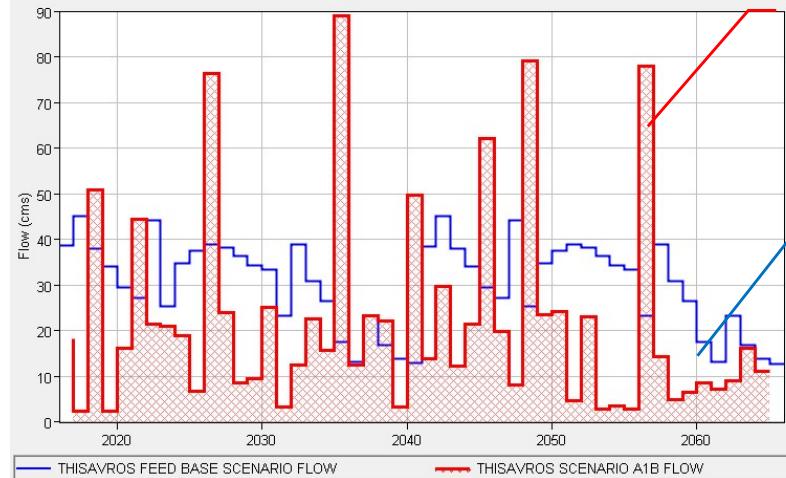
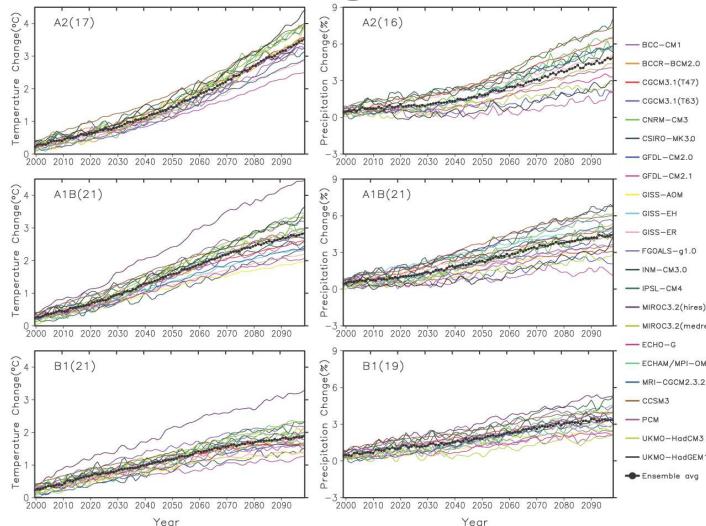
Meteorological data



Reservoir simulation



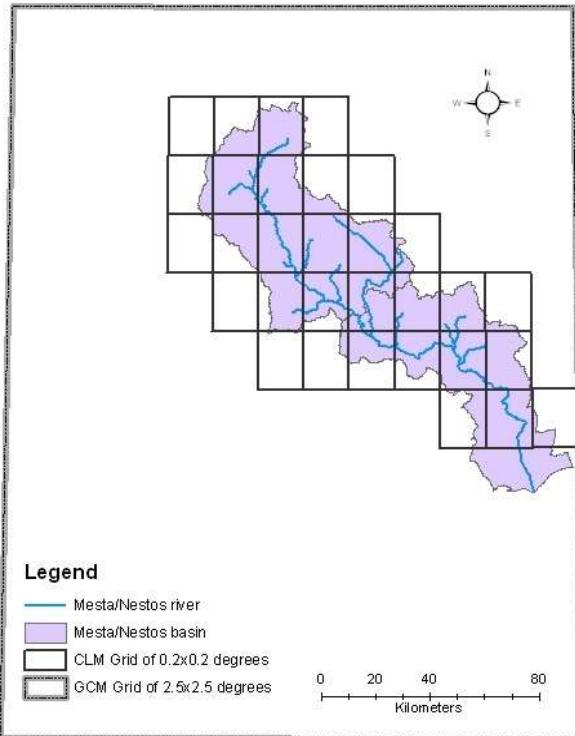
Climate change models



A1B Scenario

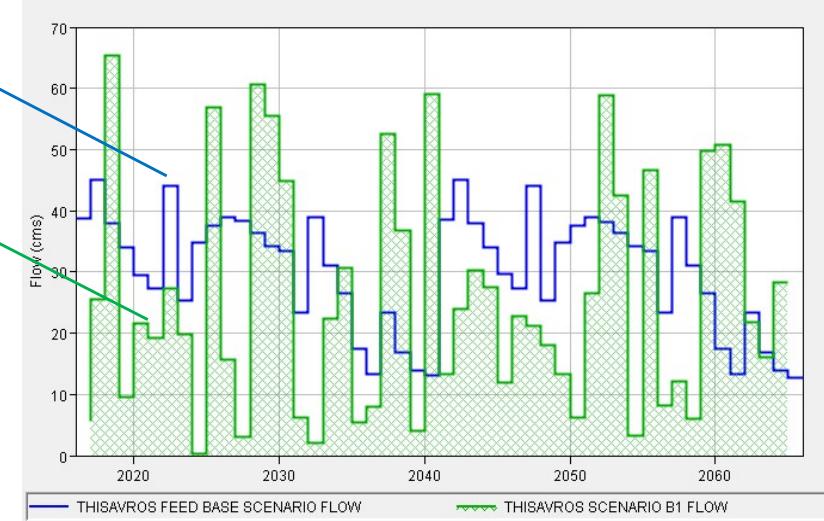
Base Scenario

Downscaling techniques



Base Scenario

B1 Scenario



	Base	A1B	B1
Mean flow (m ³ /day)	30.15	22.20	25.90

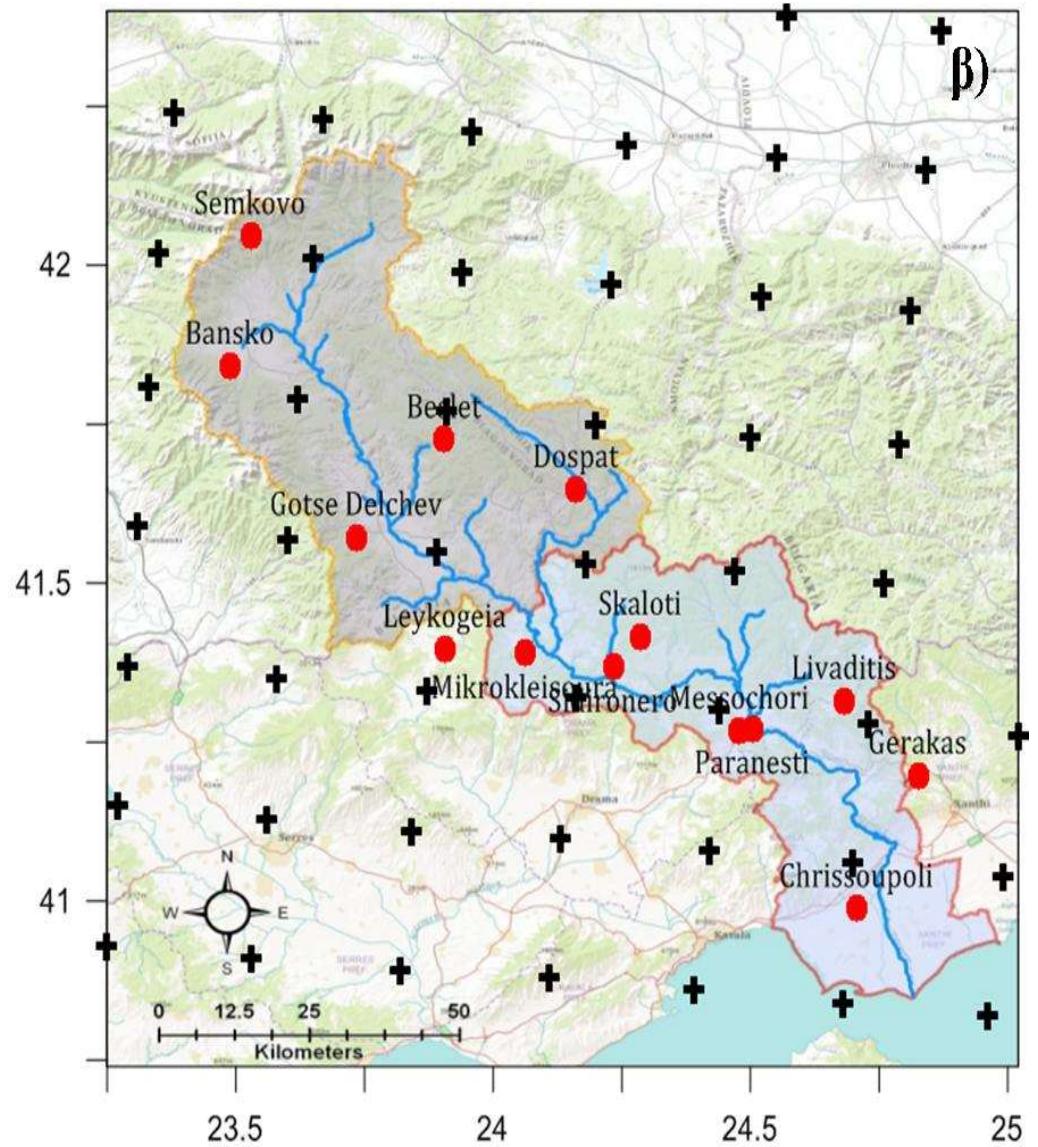
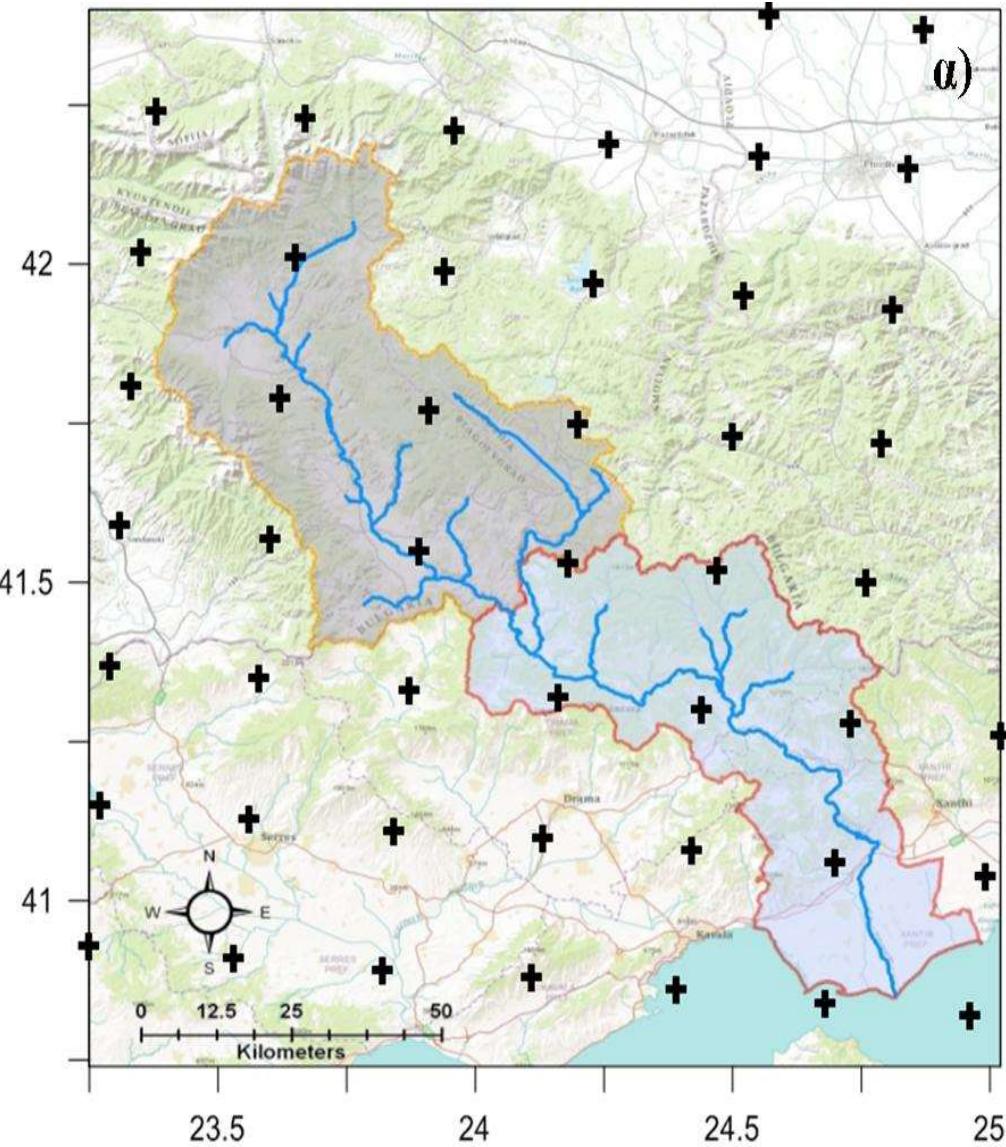
Case study: Mesta/Nestos River Basin



Climate change RCMs

MODEL	CLIMATE INSTITUTE	ABREVIATION	GCM	EMISSION SCENARIO
RCA3	Community Climate Change Consortium for Ireland	C4IRCA3	ECHAM5	A2 (- >2050)
RACMO2	Royal Netherlands Meteorological Institute	KNMI-RACMO2	ECHAM5-r3	A1B (- >2100)
RegCM3	International Centre for Theoretical Physics	ICTP-RegCM3	ECHAM5-r3	A1B (- >2100)
REMO	Max Planck Institute for Meteorology	MPI-M-REMO	ECHAM5-r3	A1B (- >2100)
CLM	Max Planck Institute for Meteorology	RCM-CLM	ECHAM5	B1 (- > 2100)

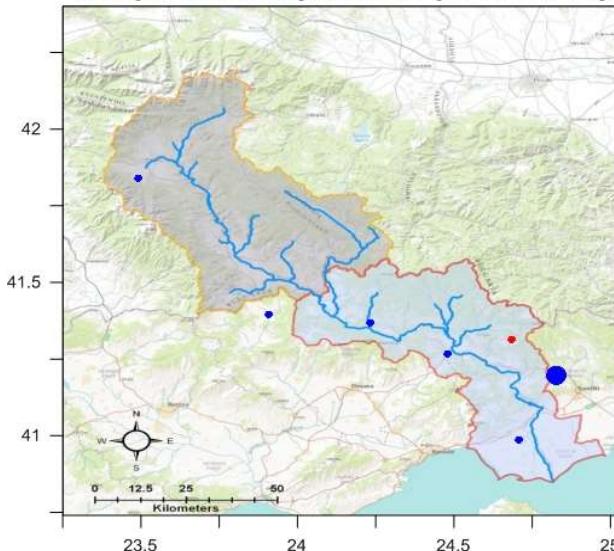
Climate change potential impacts to HPP



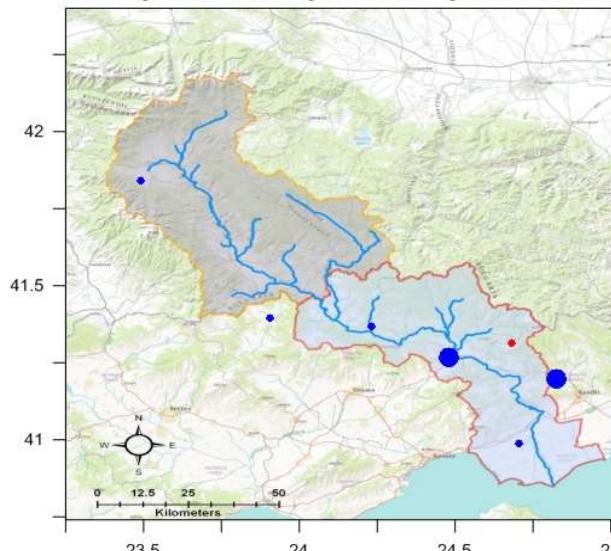
Hindcast: Temperature

Annual Temperature Differences

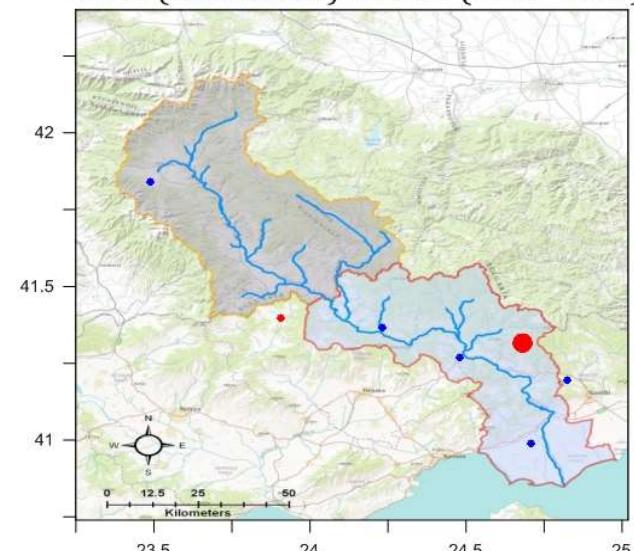
C4I (1981-2000)-Station (1981-2000)



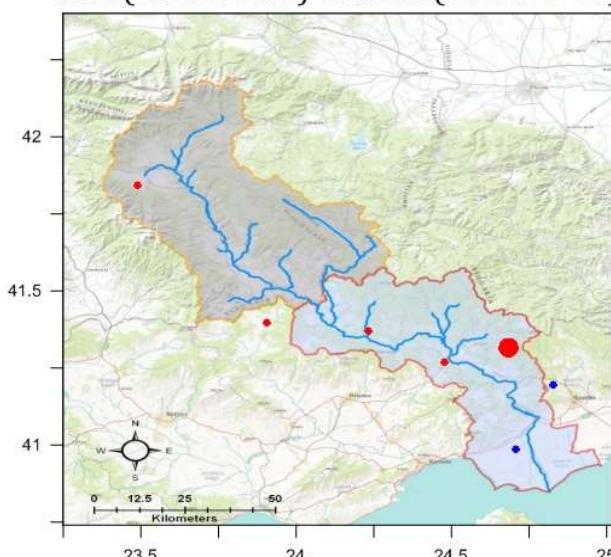
ICTP (1981-2000)-Station (1981-2000)



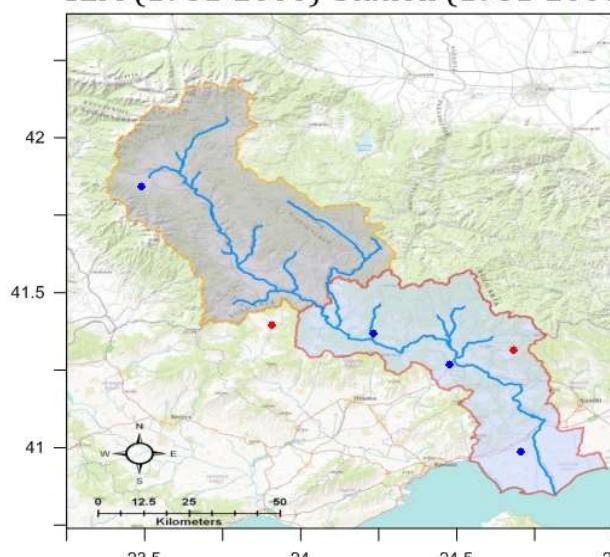
KNMI (1981-2000)-Station (1981-2000)



MPI (1981-2000)-Station (1981-2000)



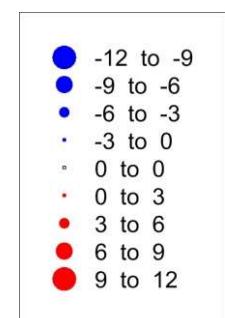
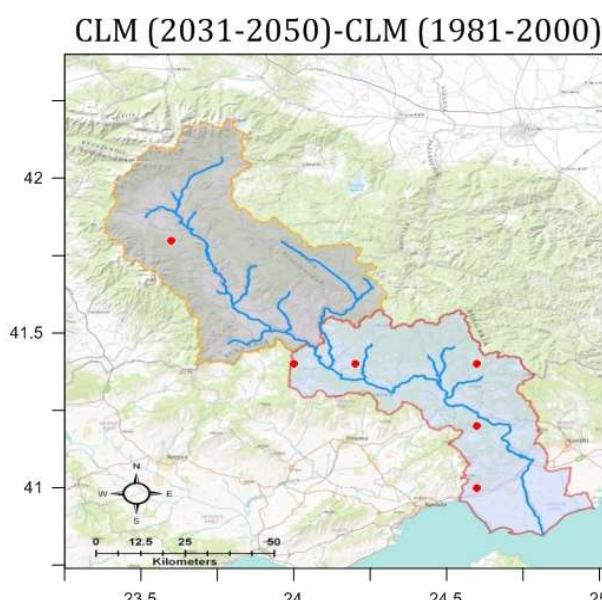
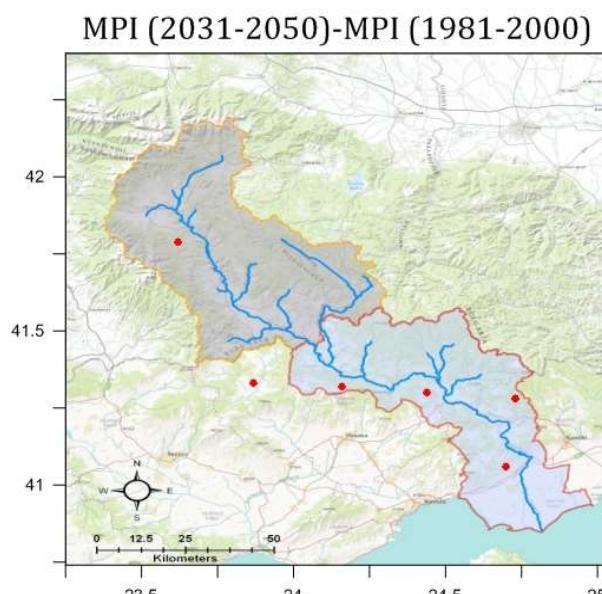
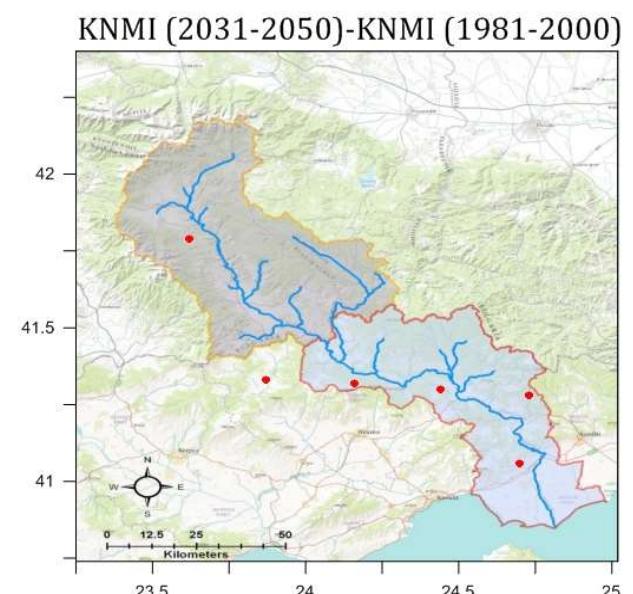
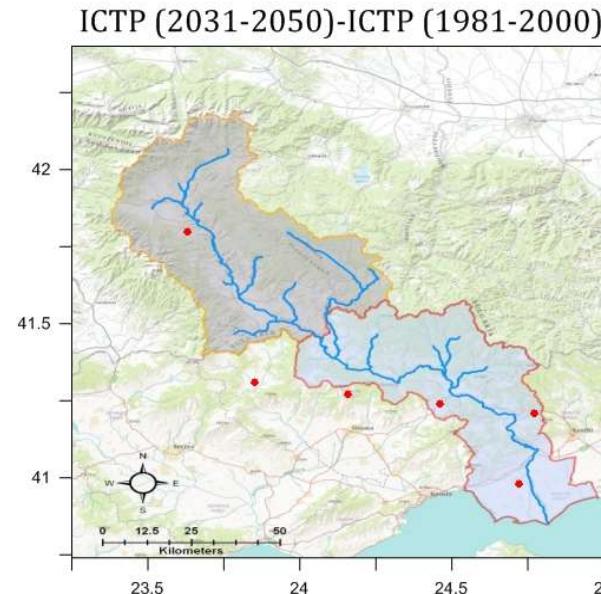
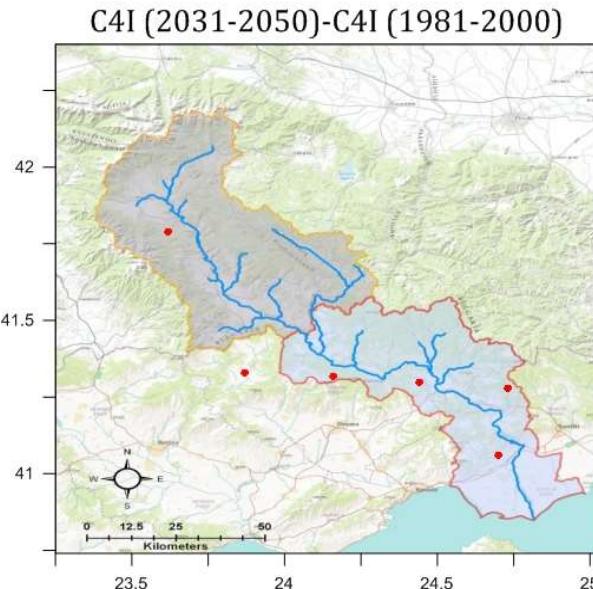
CLM (1981-2000)-Station (1981-2000)



- -12 to -9
- -9 to -6
- -6 to -3
- -3 to 0
- 0 to 0
- 0 to 3
- 3 to 6
- 6 to 9
- 9 to 12

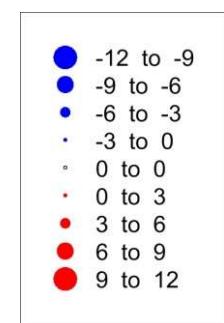
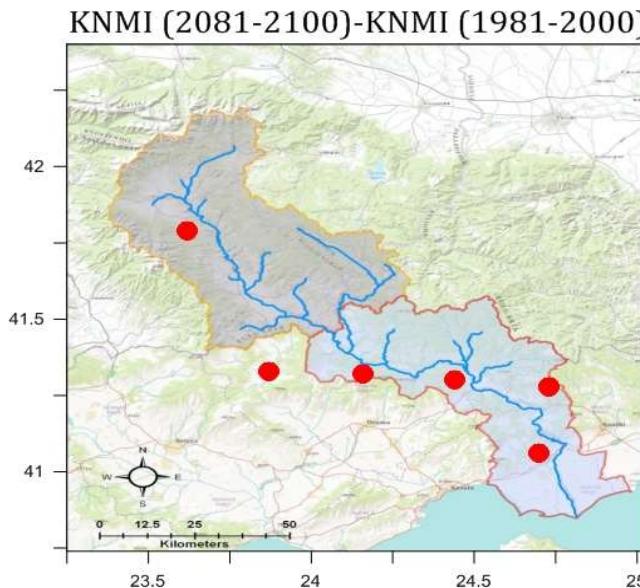
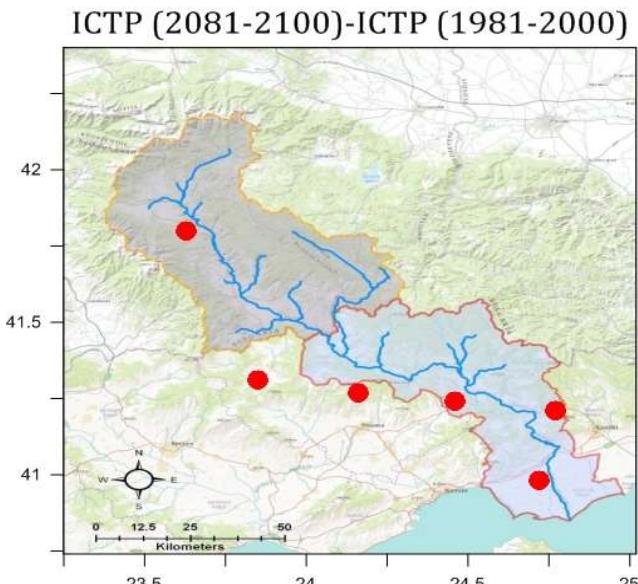
Temperature differences (1)

Annual Temperature Differences



Temperature differences (2)

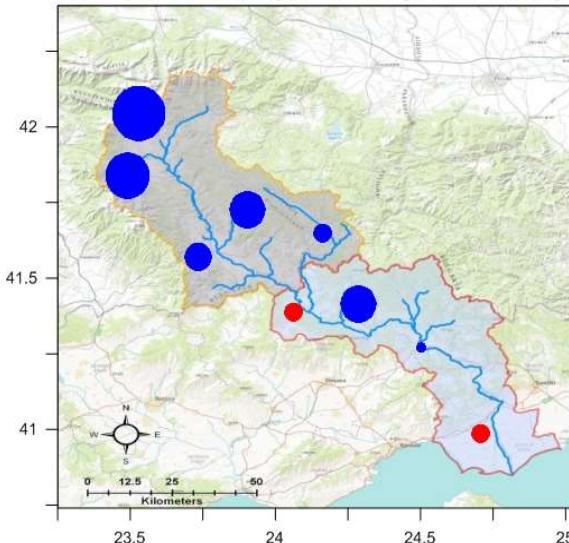
Annual Temperature Differences



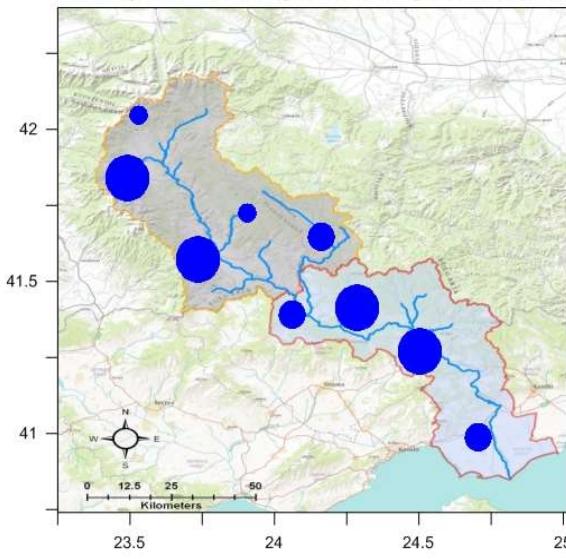
Hindcast: Precipitation

Annual Precipitation Differences

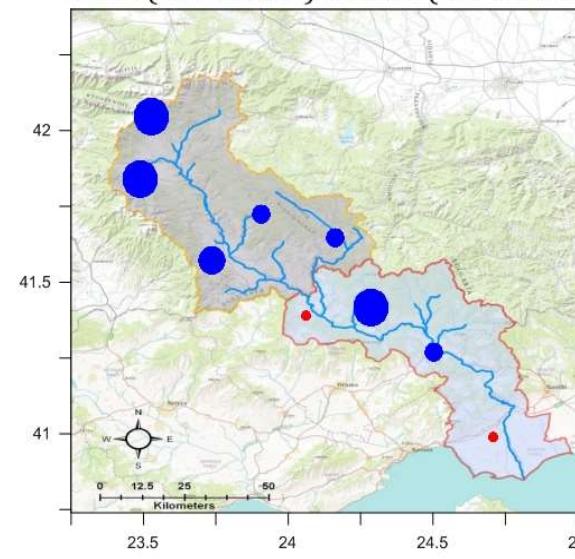
C4I (1970-1989)-Station (1970-1989)



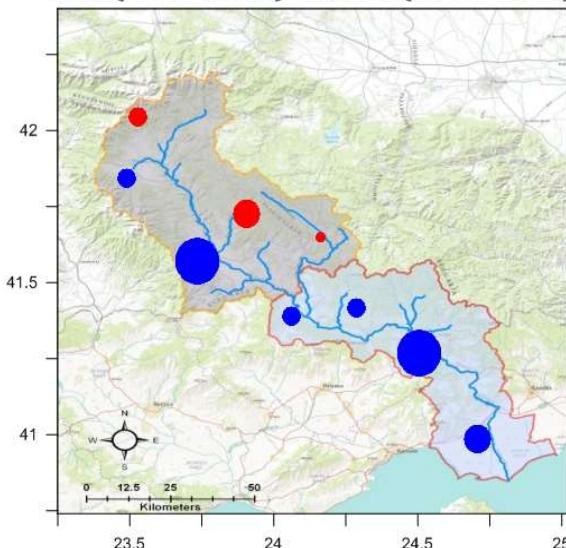
ICTP (1970-1989)-Station (1970-1989)



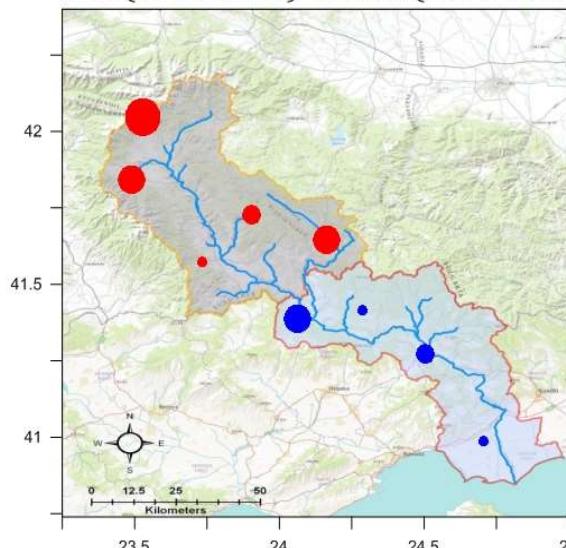
KNMI (1970-1989)-Station (1970-1989)



MPI (1970-1989)-Station (1970-1989)



CLM (1970-1989)-Station (1970-1989)

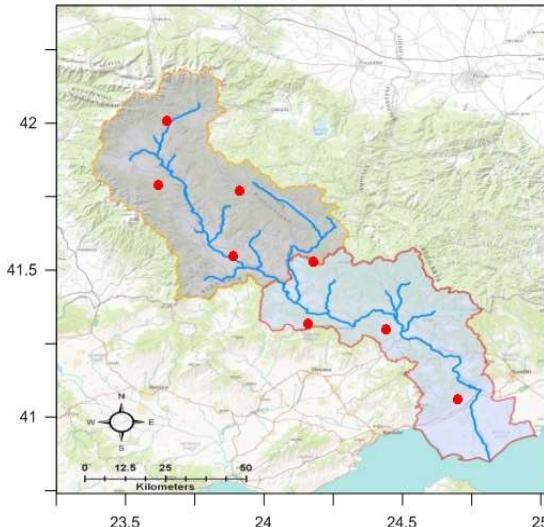


- 500 to 1000
- 400 to 500
- 300 to 400
- 200 to 300
- 100 to 200
- 0 to 100
- -100 to 0
- -200 to -100
- -300 to -200
- -400 to -300
- -500 to -400
- -1000 to -500

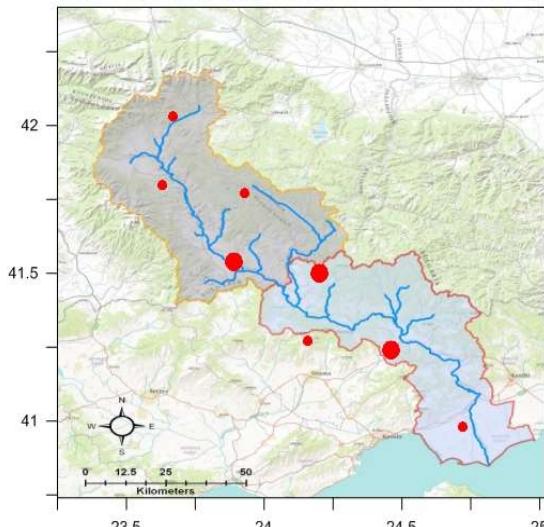
Precipitation differences (1)

Annual Precipitation Differences

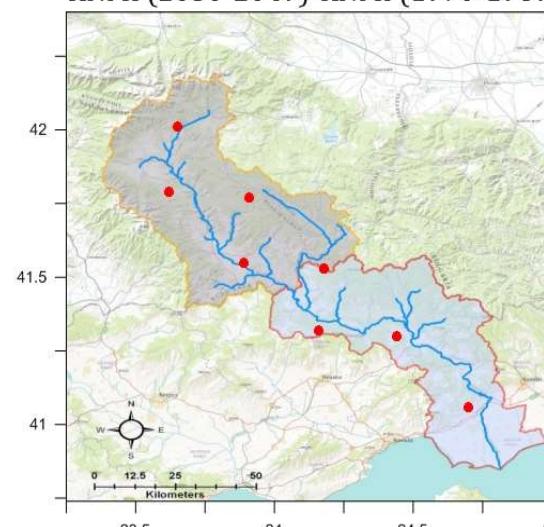
C4I (2030-2049)-C4I (1970-1989)



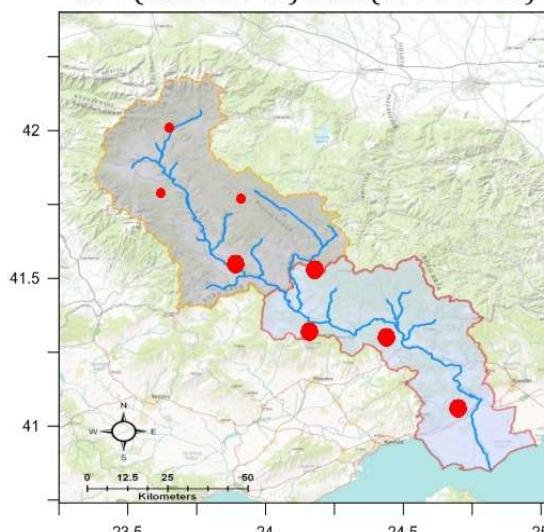
ICTP (2030-2049)-ICTP (1970-1989)



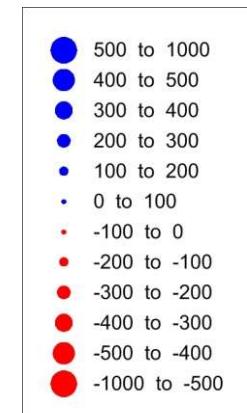
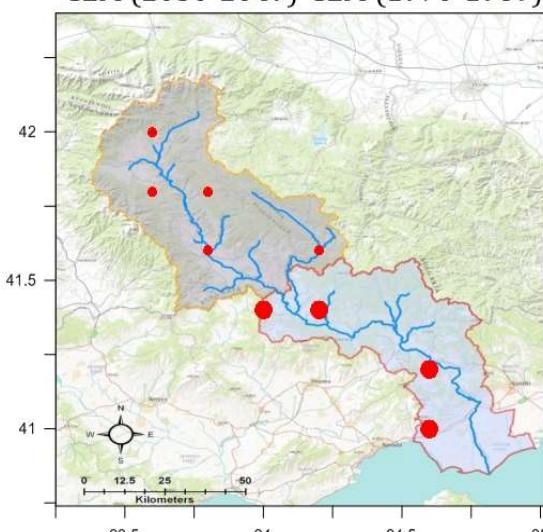
KNMI (2030-2049)-KNMI (1970-1989)



MPI (2030-2049)-MPI (1970-1989)

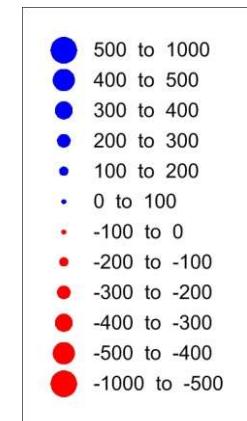
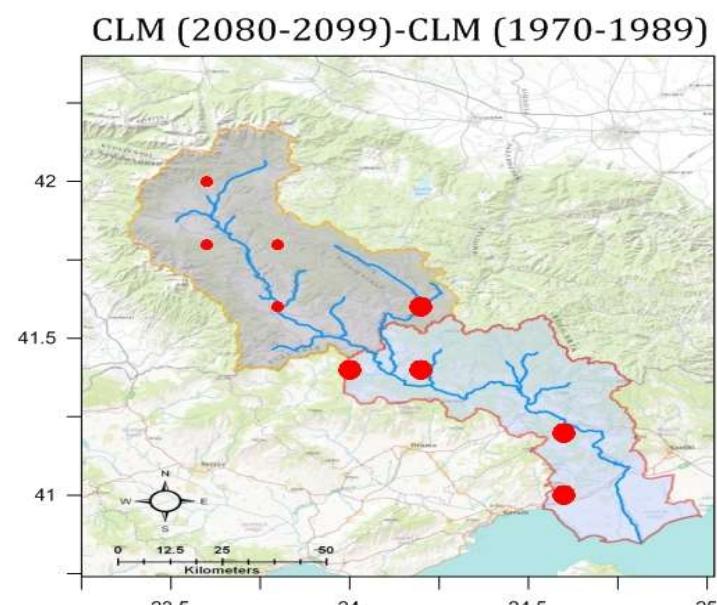
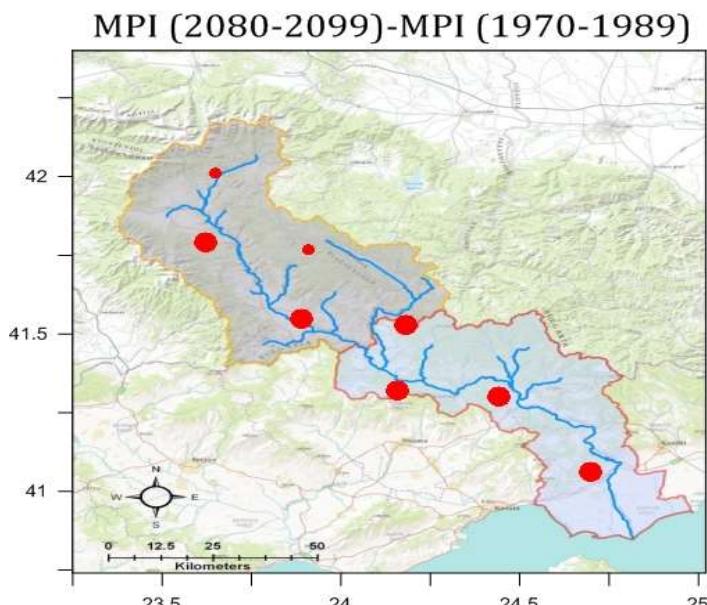
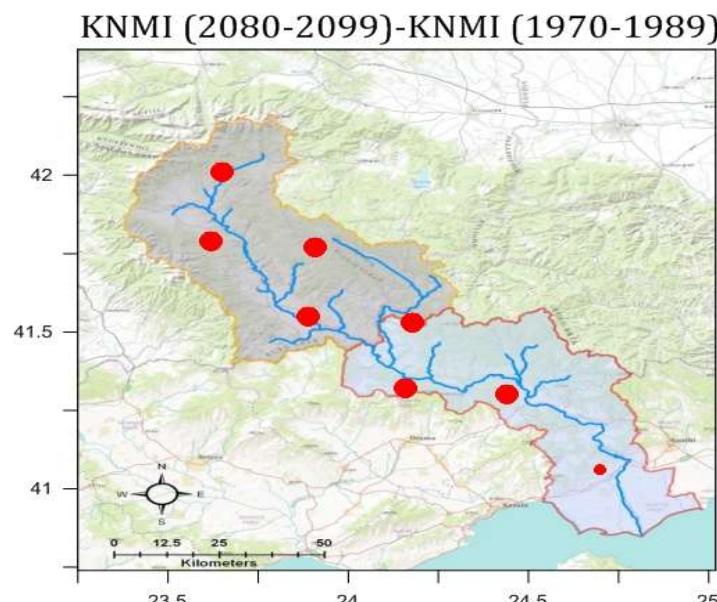
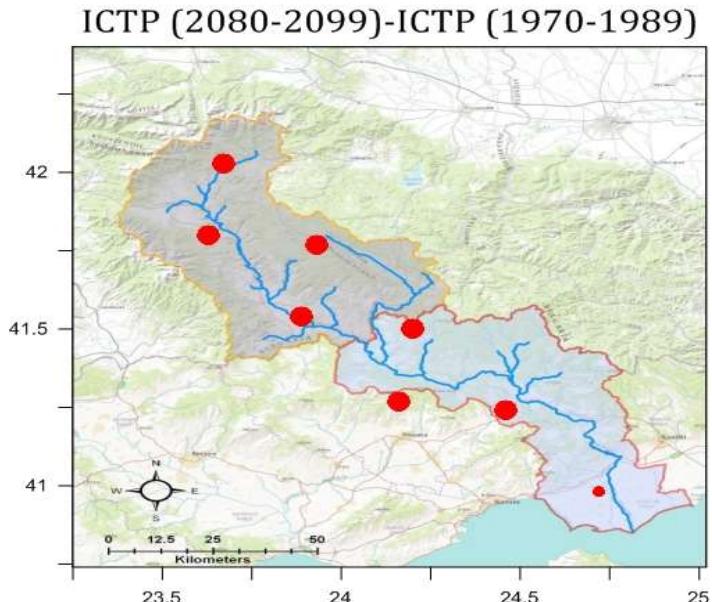


CLM (2030-2049)-CLM (1970-1989)



Precipitation differences (2)

Annual Precipitation Differences



Evapotranspiration variations

Annual Evapotranspiration Differences

